

# 柴北缘赛什腾中二叠世辉长岩成因及其对宗务隆洋盆俯冲时限的制约

庄玉军, 幸平阳, 高永伟, 彭璇, 何世平, 李普涛

(中国地质调查局 西安地质调查中心, 国土资源部岩浆作用成矿与找矿重点实验室, 陕西 西安 710054)

**摘要:** 对侵位于柴北缘构造带西段达肯大坂岩群中的赛什腾钙碱性系列辉长岩进行了地球化学、年代学研究, 结果显示, 该辉长岩具有低  $\text{SiO}_2$ 、高  $\text{Al}_2\text{O}_3$ 、富  $\text{MgO}$  及  $\text{FeO}^T$ 、贫 ALK 的特征, 且富集 LILE 元素、亏损 HFSE 元素, 稀土元素配分曲线整体表现为轻稀土元素相对富集、重稀土元素平坦且呈右倾型配分模式, 并且低 Zr 和  $\text{Zr}/\text{Y}$ , 指示赛什腾辉长岩形成于活动大陆边缘岛弧环境。锆石 LA-ICP-MS U-Pb 结晶年龄为  $271 \pm 3$  Ma。赛什腾辉长岩原岩岩浆是地幔尖晶石二辉橄榄岩部分熔融的产物, 在演化过程中经历了弱的斜长石以及较为明显的单斜辉石的分离结晶作用, 并在上升侵位过程中遭受了弱的中上地壳的同化混染作用。结合区域构造演化史以及同时代岩浆岩的年代学和地球化学特征, 认为赛什腾辉长岩是中二叠世早期宗务隆洋壳向欧龙布鲁克地块南向俯冲的产物, 且宗务隆构造带西段俯冲-碰撞事件明显早于东段, 而东西两段俯冲-碰撞时限的差异表明宗务隆洋盆存在西早东晚的“剪刀式闭合”的可能。

**关键词:** 辉长岩; 岩石地球化学; 年代学; 岩石成因; 柴北缘

中图分类号: P588.12<sup>+4</sup>; P597<sup>+3</sup>

文献标识码: A

文章编号: 1000-6524(2020)06-0718-17

## Petrogenesis of Middle Permian gabbro in Saishiteng Mountain of the northern Qaidam Basin and its constraint to the time of Zongwulong Ocean subduction

ZHUANG Yu-jun, GU Ping-yang, GAO Yong-wei, PENG Xuan, HE Shi-ping and LI Pu-tao

(Key Laboratory for the Study of Focused Magmatism and Giant Ore Deposits, MLR, Xi'an Geological Survey Center, Xi'an 710054, China)

**Abstract:** The gabbros were emplaced in the Dakendaban Group in the northwest of Saishiteng Mountain, west section of the tectonic belt of northern margin of Qaidam Basin. They are characterized by low  $\text{SiO}_2$ , high  $\text{Al}_2\text{O}_3$ , rich  $\text{MgO}$  and  $\text{FeO}^T$  but poor ALK, thus belonging to subalkalic calc-alkaline series. Meanwhile, The rocks are enriched in LILE and depleted in HFSE. The chondrite-normalized REE patterns show right-inclined shape with LREE enrichment and flat HREE. The rocks are characterized by low Zr and  $\text{Zr}/\text{Y}$  ratio. The authors consider that the gabbros emplaced in the Saishiteng Mountain were formed in the island-arc environment of an active continental margin. LA-ICP-MS zircon U-Pb dating of gabbros yielded the formation time of  $271 \pm 3$  Ma. Comprehensive analysis shows that the original rocks of gabbros can be considered as a product of partial melting of spinel lherzolite in the lithospheric mantle, and the magma might have undergone fractional crystallization of strong clinopyroxene and weak plagioclase during its evolution, which however was slightly assimilated and contaminated by mid-upper crust materials in the emplacement process. In combination with the tectonic evolution of the tectonic belt of northern margin of Qaidam Basin and the geochronological and geochemical characteristics of contemporary intrusive rocks,

收稿日期: 2020-07-27; 接受日期: 2020-10-09; 编辑: 尹淑萍

基金项目: 中国地质调查局资助项目(DD20190069); 国家自然科学基金资助项目(41602096)

作者简介: 庄玉军(1989-), 男, 汉族, 硕士, 工程师, 从事区域地质调查工作, E-mail: 179966387@qq.com。

it is concluded that the gabbros were produced in the process when the southward subduction of Zongwulong oceanic crust toward Oulongbuluke block in early Middle Permian. The authors hold that the subduction-collision event of the western part of Zongwulong structural belt was obviously earlier than that of its eastern segment, and the difference of the subduction-collision time shows that the Zongwulong Ocean were probably characterized by scissor-like closure early in the west and late in the east.

**Key words:** gabbro; geochemistry; zircon U-Pb geochronology; petrogenesis; northern margin of the Qaidam

**Fund support:** China Geological Survey Project (DD20190069); National Natural Science Foundation of China (41602096)

柴达木盆地北缘(柴北缘)构造带位于青藏高原北部,处于南祁连地块与柴达木地块的拼合部位,南东和北西两端分别被哇洪山断裂和阿尔金断裂切断(王惠初等,2005;宋述光等,2009),是一个构造复杂、物质组成多样、时间跨度大的多单元复合构造带(郭安林等,2009),以鱼卡断裂和宗务隆-青海南山断裂为界,由南向北可将柴北缘构造带划分为柴北缘早古生代俯冲带、欧龙布鲁克地块以及宗务隆构造带3个次一级构造单元(潘桂棠等,2002;王惠初等,2005)。20世纪90年代以来,柴北缘因发现早古生代大陆深俯冲的高压-超高压变质岩(杨经绥等,1998;宋述光等,2001;孟繁聪等,2003;陈丹玲等,2007)而引起了国内外地学工作者的广泛关注。前人对区内早古生代地球动力学背景和构造演化等方面进行了大量的研究工作,认为柴北缘地区在早古生代(晚寒武世)-晚古生代早期(泥盆纪-早石炭世早期)经历了洋壳俯冲(史仁灯等,2003;王惠初等,2005;高晓峰等,2011;宋述光等,2015)、(吴才来等,2007;周宾等,2013)、碰撞后板块折返(Zhang et al., 2011;邱士东等,2015)以及后造山陆内伸展(吴才来等,2007;周宾等,2013;庄玉军等,2019)的完整的造山旋回。早石炭世之后,柴北缘地区构造应力场逐渐由拉张转为收缩,辛后田等(2006)认为区内这种应力转换是由早石炭世-早二叠世邻区巴颜喀拉洋的扩张所致,吴才来等(2008)认为与古特提斯洋的关闭有关,而郭安林等(2009)则认为是柴北缘构造带北部的宗务隆构造带由洋盆发育向洋壳俯冲的转变引起的。此外,关于晚古生代晚期-早中生代柴北缘所处的构造环境也存在较大的争议。吴才来等(2008)认为柴北缘西段三岔沟一带花岗岩的形成表明中二叠世柴北缘处于加里东造山后的陆内俯冲环境下;董增产等(2014a,

2014b, 2015a, 2015b)以及辜平阳等(2016<sup>①</sup>, 2018)通过研究认为欧龙布鲁克地块西北缘晚二叠世花岗岩形成于火山弧构造环境,表明柴北缘晚二叠世处于洋陆俯冲的构造演化阶段,但吴锁平(2008)、邱士东等(2015)、高万里等(2019)则认为柴北缘在中二叠世已经处于洋陆俯冲阶段。杨明慧等(2002)认为小寨什腾山一带的中三叠世花岗岩体具S型花岗岩特征,进而认为中三叠世柴北缘已处于俯冲结束后的同碰撞阶段,而强娟(2008)、彭渊等(2016)、王苏里等(2016)则认为中三叠世柴北缘的洋壳俯冲作用仍未结束。由此可见,柴北缘晚古生代晚期-早中生代构造演化背景争议较大,尚需要进一步研究。

综上可知,大多学者对柴北缘晚古生代的研究对象多为中酸性侵入岩,而中酸性侵入岩的地球化学构造背景往往存在多解性。相比而言,基性-超基性岩浆岩更能反映构造背景,所以通常被作为研究构造环境演化的重要研究对象(Zhang et al., 2009;张云等,2020),然而该地区晚古生代的基性-超基性岩很少。最近,笔者在柴北缘赛什腾山地区开展专项地质调查时,在古元古界达肯大坂岩群中识别出的中二叠世辉长岩,为探讨柴北缘地区晚古生代的构造演化提供了新的载体。本文通过对该辉长岩的地球化学、年代学研究,探讨其岩石成因及构造环境,旨在为柴北缘晚古生代的构造格局与演化提供新的约束。

## 1 区域地质概况

研究区位于柴北缘构造带西段赛什腾山西北部,区域上出露的地层主要为古元古界达肯大坂岩群(Pt<sub>1</sub>DK.)以及下古生界滩间山群(ε-OT<sub>1</sub>)(图1a)。其中,达肯大坂岩群是一套原岩为火山-碎屑

<sup>①</sup> 辜平阳,等. 2016. 青海阿尔金1:5万打柴沟等6幅区调报告.

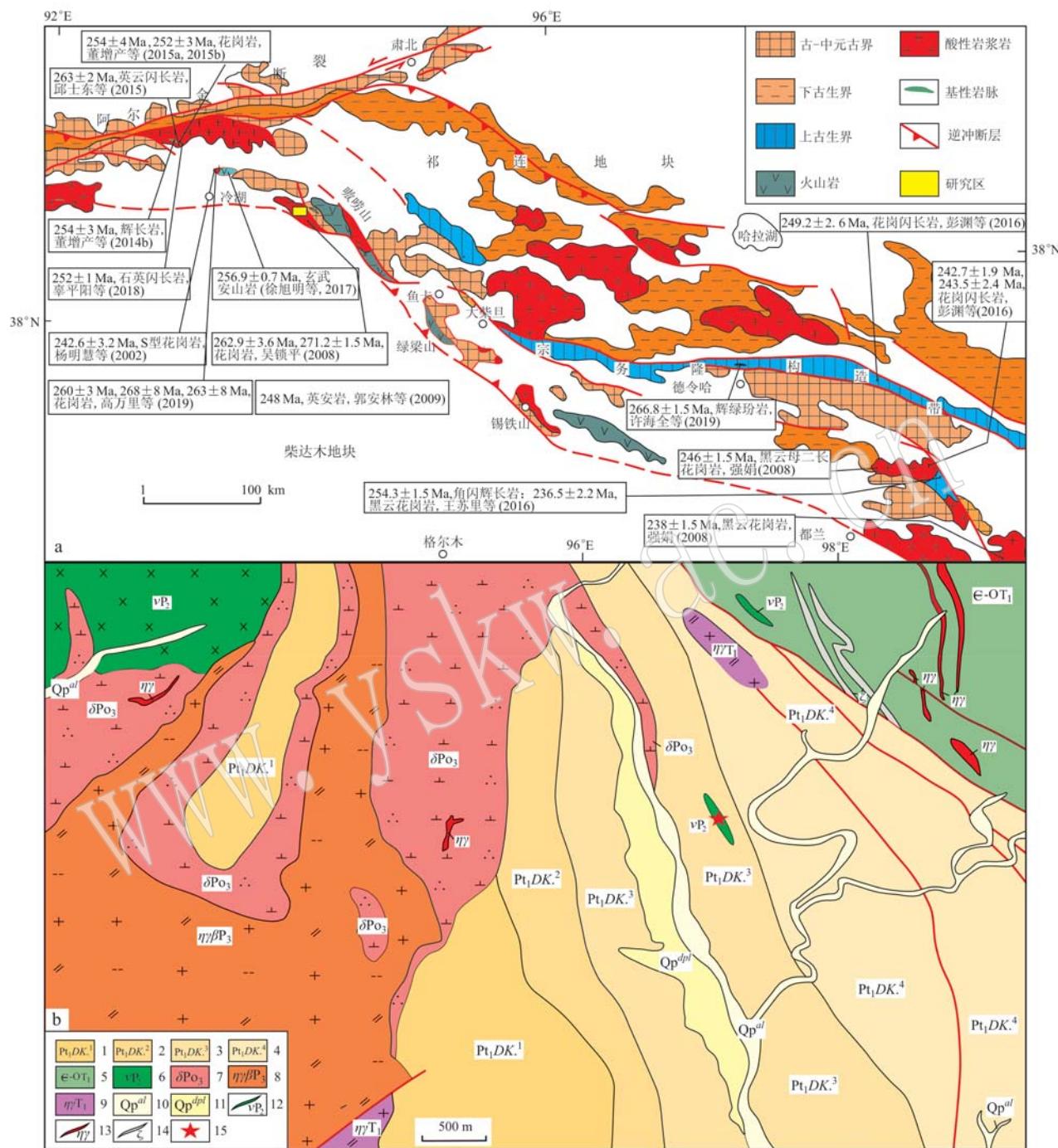


图1 柴北缘地质简图[a, 据杨经绥等(2001)修改]及研究区地质图(b)

Fig. 1 Sketch geological map of the northern Qaidam Basin (a, after Yang Jingsui et al., 2001) and the geological map of the study areas (b)

- 1—达肯大坂岩群第一岩性段; 2—达肯大坂岩群第二岩性段; 3—达肯大坂岩群第三岩性段; 4—达肯大坂岩群第四岩性段; 5—滩间山群;  
 6—辉长岩体; 7—石英闪长岩体; 8—黑云二长花岗岩体; 9—二长花岗岩体; 10—冲积物; 11—坡洪积物; 12—辉长岩脉;  
 13—二长花岗岩脉; 14—英安岩; 15—采样点  
 1—1st Member of Dakendaban Group; 2—2nd Member of Dakendaban Group; 3—3rd Member of Dakendaban Group; 4—4th Member of Dakendaban Group; 5—Tanjianshan Group; 6—gabbro massif; 7—quartz diorite body; 8—biotite adamellite rock body; 9—adamellite rock body; 10—alluvium; 11—slope-proluvium; 12—gabbro dike; 13—adamellite vein; 14—dacite; 15—sampling point

岩系并经历了中高级变质的副变质岩;滩间山群下部以中基性海相火山岩为主,但徐旭明等(2017)在

其中的玄武安山岩中获得  $256.9 \pm 0.7$  Ma 的锆石 LA-ICP-MS(U-Pb) 年龄。区内基性、中性、酸性岩浆岩

均有出露,且以中酸性岩为主,主要为英云闪长岩、花岗闪长岩、二长花岗岩以及花岗伟晶岩等,时代有奥陶纪、志留纪、石炭纪、二叠纪及三叠纪,并以二叠纪、奥陶纪最为发育;基性岩(脉)则以辉长岩为主,侵入于达肯大坂岩群、滩间山群中,并被晚二叠世石英二长岩侵入。本文研究的辉长岩脉产状近直立( $75^{\circ} \sim 85^{\circ}$ ),呈北西西向( $290^{\circ}$ )延伸400余米,最宽处可达70余米,沿片理、片麻理侵入达肯大坂岩群黑云斜长片麻岩及二云母石英片岩中( $Pt_1DK.3$ )中,并被后期北西西向的花岗细晶岩脉侵入。

## 2 岩石学特征

用于同位素及地球化学研究的辉长岩样品(编号为PM002-62-1)采自冷湖镇东( $105^{\circ}$ ) $53\text{ km}$ 处,地理坐标 $93^{\circ}55'35''\text{E}, 38^{\circ}36'40''\text{N}$ 。样品风化色呈灰绿色、灰黑色,具块状构造,变余辉长结构(图2)。岩

石主要由斜长石( $\sim 45\%$ )、辉石( $\sim 40\%$ )、角闪石( $\sim 10\%$ )、钛铁矿( $3\%$ )以及少量榍石( $<1\%$ )等组成。斜长石与辉石晶体均呈自形、半自形粒状,二者形态与粒径大小相近,粒径大小一般在 $2 \sim 6.6\text{ mm}$ 之间,其中斜长石晶体普遍出现明显的绢云母化、黝帘石化,多数辉石已次生蚀变被角闪石交代。钛铁矿晶体多呈板状,粒径大小在 $0.15 \sim 1.2\text{ mm}$ 之间,有时被榍石交代。

## 3 分析方法

样品的主微量及稀土元素测试分析在中国地质调查局西安地质调查中心实验测试中心完成,主量元素采用SX45型X荧光光谱仪(XRF)进行分析,其中FeO含量通过湿化学方法测定采用,分析误差小于1%;微量元素和稀土元素利用SX50型电感耦合等离子体光谱仪(ICP-MS)进行测定,分析误差为

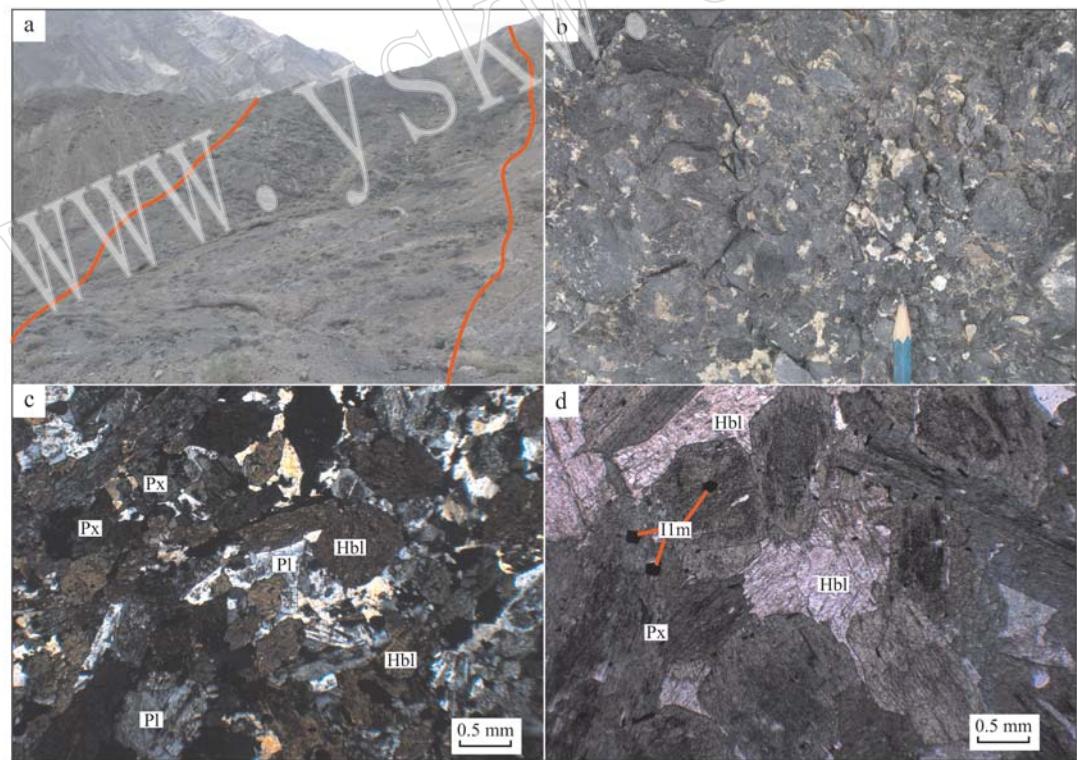


图2 赛什腾辉长岩宏观产出及显微镜下特征

Fig. 2 Macroscopic and microscopic characteristics for gabbro of Saishiteng Mountain

a—辉长岩宏观产出特征; b—辉长岩野外露头; c—变余辉长结构(+); d—变余辉长结构(-); Hbl—角闪石; Px—辉石; Pl—斜长石; Ilm—钛铁矿

a—macroscopic production characteristics of gabbro; b—field outcrop of gabbro; c—palimpsest gabbro texture (+); d—palimpsest gabbro texture (-); Hbl—hornblende; Px—pyroxene; Pl—plagioclase; Ilm—ilmenite

5%~10%。样品锆石挑选由河北廊坊诚信地质服务有限公司完成,锆石的制靶及反射光阴极发光照相在陕西爱思拓普测试技术有限公司完成,测试点的选取首先根据锆石反射光和透射光照片进行初选,再与CL图像反复对比,力求避开内部裂隙和包裹体,以获得较准确的年龄信息。LA-ICP-MS锆石微区U-Pb年龄测定在自然资源部岩浆作用成矿与找矿重点实验室完成,采用193nmArF准分子(excimer)激光器的Geo Las200M剥蚀系统,ICP-MS为Agilent 7700,激光束斑直径24 μm,以GJ-1为同位素监控标样,91500为年龄标定标样,NIST610为元素含量标样进行校正,普通铅校正依据实测<sup>204</sup>Pb进行校正。

采用Glitter(ver4.0, Macquarie University)程序对锆石的同位素比值及元素含量进行计算,并按照Andersen Tom的方法(Andersen, 2002),用LAM-ICPMS Common Lead Correction(ver3.15)对其进行普通铅校正,年龄计算及谐和图采用Isoplot(ver3.0)完成(Ludwig, 2003)。

## 4 分析结果

### 4.1 全岩地球化学特征

在主量元素分析表(表1)中可以看出,赛什腾

**表1 辉长岩主量元素( $w_B/\%$ )、微量元素及稀土元素( $w_B/10^{-6}$ )含量分析结果**

**Table 1 Major elements ( $w_B/\%$ ), trace elements and REE ( $w_B/10^{-6}$ ) compositions of gabbro**

样号	PM002-62-1-1	PM002-62-1-2	PM002-62-1-3	PM002-62-1-4	PM002-62-1-5	PM002-62-1-6
SiO <sub>2</sub>	45.34	47.69	46.26	45.92	47.05	46.76
TiO <sub>2</sub>	1.35	1.33	1.59	1.61	1.51	1.61
Al <sub>2</sub> O <sub>3</sub>	19.45	19.00	17.92	17.68	18.15	17.31
Fe <sub>2</sub> O <sub>3</sub>	5.60	3.86	3.65	3.69	3.98	3.96
FeO	6.45	6.11	6.78	6.03	6.40	6.44
MnO	0.15	0.13	0.14	0.13	0.13	0.15
MgO	5.90	5.80	7.51	8.48	6.95	7.81
CaO	9.93	9.88	9.86	10.65	9.72	9.90
Na <sub>2</sub> O	2.16	2.35	2.06	2.22	2.21	2.45
K <sub>2</sub> O	1.15	1.28	1.29	0.91	1.21	1.05
P <sub>2</sub> O <sub>5</sub>	0.18	0.17	0.17	0.10	0.16	0.17
LOI	2.34	2.40	2.77	2.58	2.53	2.39
Total	100.00	100.00	100.00	100.00	100.00	100.00
FeO <sup>T</sup>	11.49	9.58	10.06	9.35	9.98	10.00
Mg <sup>#</sup>	47.82	51.93	57.12	61.82	55.42	58.22
m/f	0.90	1.07	1.31	1.60	1.23	1.37

续表1  
Continued Table 1

样号	PM002-62-1-1	PM002-62-1-2	PM002-62-1-3	PM002-62-1-4	PM002-62-1-5	PM002-62-1-6
σ	3.35	2.36	2.59	2.50	2.34	2.62
Ba	196.00	248.00	275.00	181.00	249.00	209.00
Rb	39.50	44.50	44.70	27.90	35.70	36.50
Sr	706.00	696.00	691.00	713.00	671.00	696.00
Y	19.60	19.60	19.50	19.00	17.20	23.00
Zr	32.90	57.90	45.90	36.30	38.90	59.10
Nb	5.65	6.20	6.21	5.05	5.70	7.41
Th	2.99	2.41	3.12	2.36	3.43	3.39
Ga	24.80	22.70	21.50	18.90	20.20	21.00
Ni	24.90	45.50	63.20	41.80	55.10	78.80
Cr	35.90	58.40	105.00	55.60	83.00	75.30
Hf	1.56	2.16	2.36	1.60	1.60	2.20
Sc	26.60	30.00	39.30	44.90	34.40	42.20
Ta	0.66	0.68	0.67	0.51	0.61	0.61
Co	39.70	34.10	41.00	40.10	38.50	41.50
Li	25.40	21.20	26.70	14.40	22.70	23.10
Be	0.89	1.10	0.89	0.68	0.85	0.91
V	380.00	372.00	428.00	402.00	390.00	408.00
Cs	3.35	3.66	4.35	2.74	3.77	3.98
U	0.83	1.54	0.87	0.46	0.88	0.71
Pb	9.74	10.90	11.10	7.07	10.90	9.76
Cu	68.30	68.60	28.10	45.30	44.10	30.20
Zn	99.30	96.40	89.20	67.80	81.80	90.60
Ge	1.39	1.56	1.78	1.57	1.49	1.44
La	11.00	10.60	13.40	13.10	12.00	16.40
Ce	28.20	27.20	31.10	29.00	28.40	39.60
Pr	4.09	3.99	4.32	4.03	3.92	5.34
Nd	18.80	17.80	19.30	17.80	17.00	23.10
Sm	4.69	4.30	4.54	4.38	3.99	5.33
Eu	1.52	1.21	1.29	1.31	1.16	1.48
Gd	4.46	4.12	4.24	4.36	3.92	5.17
Tb	0.72	0.69	0.68	0.66	0.65	0.82
Dy	4.06	3.90	4.05	3.98	3.61	4.77
Ho	0.81	0.77	0.82	0.80	0.72	0.92
Er	2.14	2.11	2.19	2.08	1.95	2.51
Tm	0.32	0.31	0.32	0.30	0.28	0.36
Yb	1.96	2.08	2.05	1.86	1.83	2.24
Lu	0.30	0.30	0.30	0.28	0.27	0.34
δEu	1.00	0.87	0.89	0.91	0.89	0.85
δCe	1.01	1.01	0.98	0.95	0.99	1.01
LREE	68.30	65.10	73.95	69.62	66.47	91.25
HREE	34.37	33.88	34.15	33.31	30.44	40.12
ΣREE	102.67	98.98	108.10	102.93	96.91	131.37
LREE/HREE	1.99	1.92	2.17	2.09	2.18	2.27
Nb/La	0.51	0.58	0.46	0.39	0.48	0.45
(La/Yb) <sub>N</sub>	3.78	3.44	4.41	4.75	4.42	4.94
(La/Sm) <sub>N</sub>	1.48	1.55	1.86	1.88	1.89	1.94
(Gd/Yb) <sub>N</sub>	1.84	1.60	1.67	1.89	1.73	1.86
(Th/Nb) <sub>PM</sub>	4.44	3.26	4.21	3.92	5.05	3.84
(La/Nb) <sub>PM</sub>	2.02	1.77	2.24	2.69	2.18	2.30
(Th/Ta) <sub>PM</sub>	2.18	1.72	2.26	2.23	2.73	2.66
(La/Sm) <sub>PM</sub>	1.52	1.59	1.91	1.93	1.94	1.99
(Tb/Yb) <sub>PM</sub>	1.67	1.51	1.52	1.61	1.63	1.68

辉长岩样品的烧失量较低(2.34%~2.77%),表明样品受后期低温蚀变作用及风化作用的影响较小。样品中 $\text{SiO}_2$ 含量较低,为45.34%~47.69%,相对富 $\text{MgO}$ (5.80%~8.48%)及 $\text{FeO}^T$ (9.35%~11.49%),贫 $\text{ALK}$ ( $\text{K}_2\text{O} + \text{Na}_2\text{O}$ 含量为3.13%~3.63%)和 $\text{P}_2\text{O}_5$ (0.10%~0.18%), $\text{TiO}_2$ 为中高含量(1.33%~1.61%),可能与岩石中发育钛铁矿副矿物有关, $\text{Al}_2\text{O}_3$ 含量为17.31%~19.45%,平均18.25%(>17%),具有高铝玄武质岩石的特征(桑隆康等,2012)。镁铁比值m/f为0.90~1.60,属铁质基性岩类(m/f=0.5~2.0); $\text{Mg}^\# = 47.82 \sim 61.82$ ,小于原始岩浆玄武岩 $\text{Mg}^\# = 68 \sim 78$ (李文宣等,1994)。扣除烧失量作归一化处理后对样品进行投图,在哈克图解中, $\text{MgO}$ 与 $\text{SiO}_2$ 、 $\text{Na}_2\text{O} + \text{K}_2\text{O}$ 以及 $\text{Al}_2\text{O}_3$ 存在较为明显的负相关关系(图3a、3b、3d), $\text{MgO}$ 与 $\text{CaO}$ 相关性较弱(图3c), $\text{MgO}$ 和 $\text{Mg}^\#$ 分别与 $\text{TiO}_2$ 和 $\text{CaO}/\text{Al}_2\text{O}_3$ 呈明显正相关关系(图3e、3f);在Cox等(1979)的TAS分类图上除样品PM002-62-1-1外,其它5个样品均落入亚碱性系列辉长岩区(图4a),在 $\text{FeO}^T - \text{FeO}^T/\text{MgO}$ 图解(图4b)

中大多数样品落入钙碱性系列范围内,仅一个样品落入拉斑玄武岩系列范围内,且所有样品的里特曼指数 $\sigma$ 均小于4(表1),表明该辉长岩为钙碱性辉长岩。

在稀土微量元素分析表(表1)中,辉长岩的稀土总量( $\Sigma\text{REE}$ )较低且变化范围不大,为 $96.91 \times 10^{-6} \sim 131.37 \times 10^{-6}$ , $\text{LREE} = 65.10 \times 10^{-6} \sim 91.25 \times 10^{-6}$ , $\text{HREE} = 30.44 \times 10^{-6} \sim 40.12 \times 10^{-6}$ , $\text{LREE}/\text{HREE} = 1.92 \sim 2.27$ , $(\text{La}/\text{Yb})_N = 3.44 \sim 4.94$ ,表明轻稀土相对富集,轻重稀土元素分异较为明显; $(\text{La}/\text{Sm})_N = 1.48 \sim 1.94$ , $(\text{Gd}/\text{Yb})_N = 1.60 \sim 1.89$ ,显示轻、重稀土内部分异均较弱;样品的 $\delta\text{Eu} = 0.85 \sim 1.00$ ,显示为弱的负异常或无异常。此外, $\delta\text{Ce} = 0.95 \sim 1.01$ ,波动范围较小,暗示样品具有一致的源区和相似的演化过程。在球粒陨石标准化稀土元素配分图上,各样品具有相似的稀土分布模式,均表现出LREE相对富集、HREE平坦且向右缓倾的配分模式。原始地幔标准化微量元素蛛网图显示,辉长岩明显富集Rb、Ba、Th、U、Sr等大离子亲石元素(LILE),显著亏损Nb、Ta、Zr高场强元素(HFSE),Ti为弱亏损(图5)。

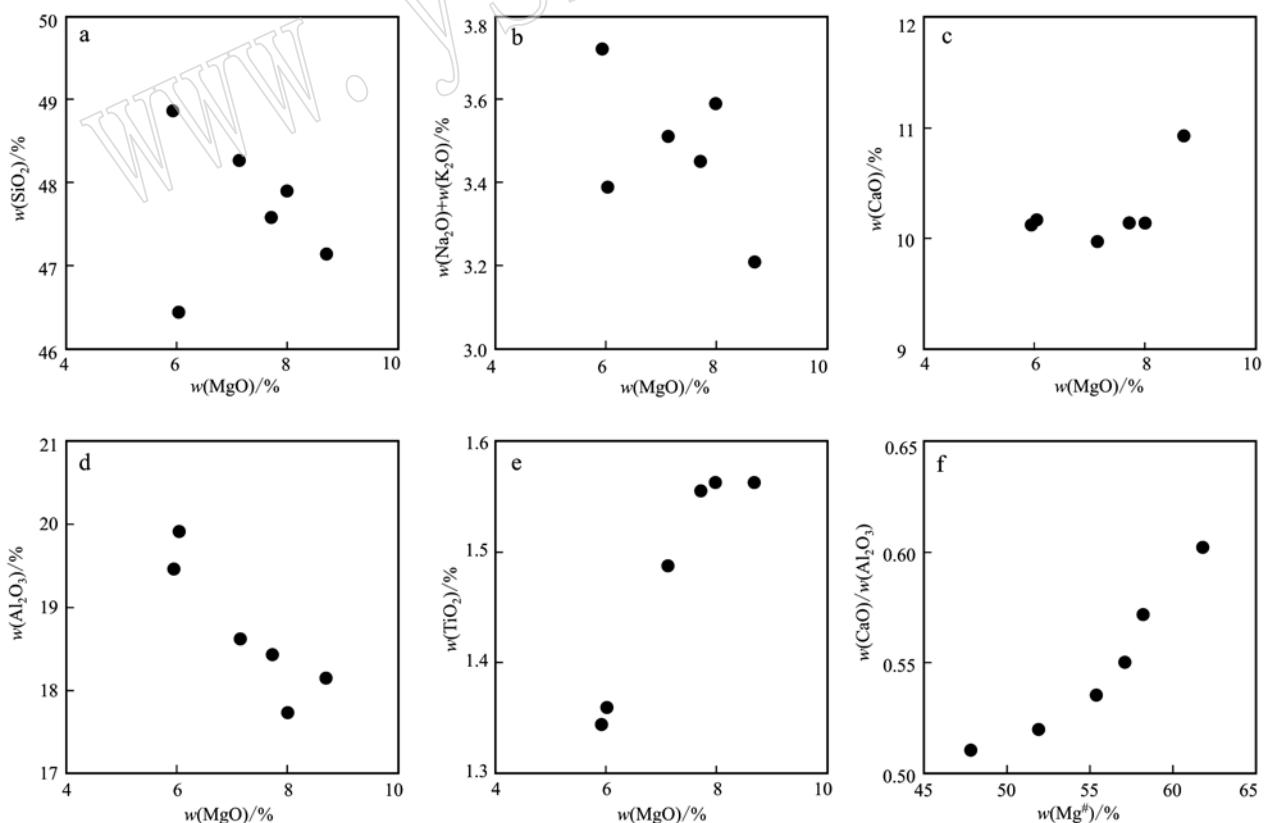


图3 赛什腾辉长岩Harker图解

Fig. 3 Harker diagrams for gabbro of Saishiteng Mountain

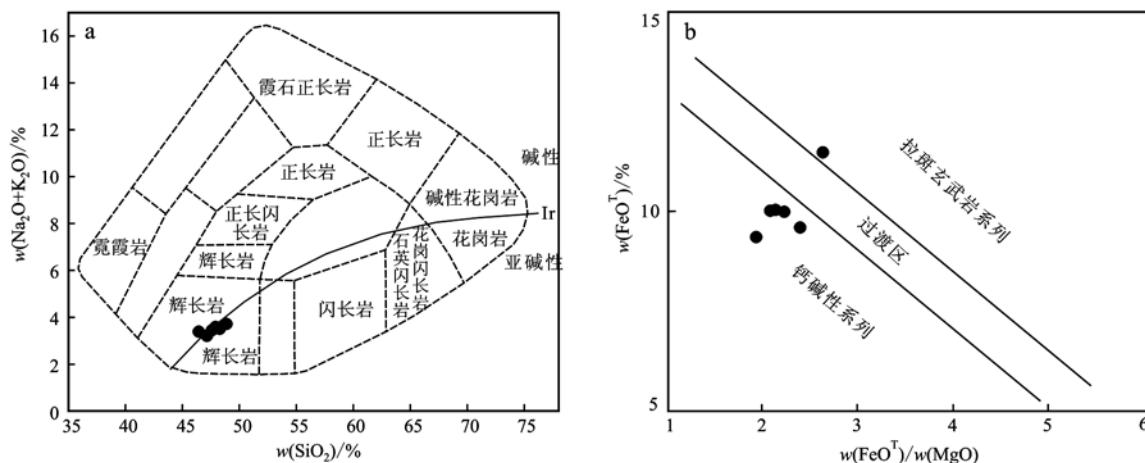


图 4 赛什腾辉长岩 TAS 分类图(a, 底图据 Wilson, 1989)及  $\text{FeO}^T - \text{FeO}^T/\text{MgO}$  图解(b, 据 Miyashiro, 1974)

Fig. 4 TAS diagram (a, after Wilson, 1989) and  $\text{FeO}^T$  versus  $\text{FeO}^T/\text{MgO}$  diagram (b, after Miyashiro, 1974) for gabbro of Saishiteng Mountain

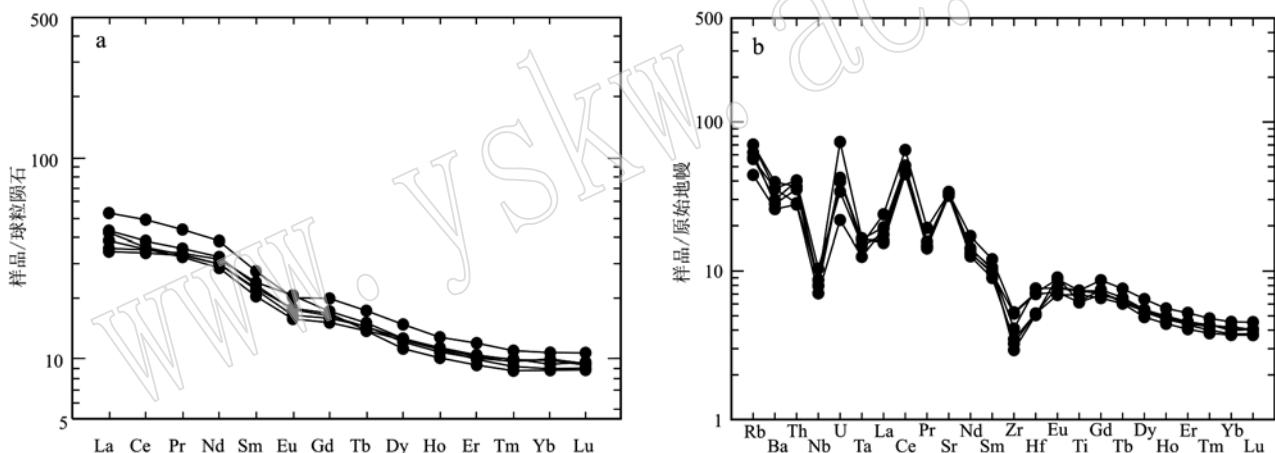


图 5 辉长岩稀土元素球粒陨石标准化图解(a)和微量元素原始地幔标准化蛛网图(b)(球粒陨石标准化值及原始地幔标准化值据 Sun and McDonough, 1989)

Fig. 5 Chondrite-normalized REE patterns (a) and primitive-mantle normalized spidergram (b) of gabbro (chondrite-normalized values and primitive mantle-normalized values after Sun and McDonough, 1989)

## 4.2 成岩时代

赛什腾辉长岩的锆石阴极发光(CL)图像见图6,锆石的U-Pb同位素比值和表面年龄测试数据列于表2。大多数锆石镜下呈无色透明,短柱状、长柱状晶形明显(长 $100\sim180\text{ }\mu\text{m}$ ),长宽比为 $1:1\sim2:1$ ,CL照片显示该类锆石发育宽缓或条带状振荡环带结构(10、12、33、34号等),个别锆石具有因后期蚀变或变质产生的浅色增生边(5号)。锆石中Th和U含量分别为 $38\times10^{-6}\sim828\times10^{-6}$ 和 $44\times10^{-6}\sim664\times10^{-6}$ ,Th/U值 $0.48\sim1.68$ ,Th和U之间的相关系数为0.90(图略)。锆石的稀土元素球粒陨石

标准化配分模式显示锆石显著富集 HREE，并具有明显的正  $\delta\text{Ce}$  异常和负  $\delta\text{Eu}$  异常（图 7a），结合锆石存在生长韵律环带、较高的 Th/U 值及 Th、U 相关系数，表明这些锆石为典型的基性岩浆锆石成因（Hoskin and Black, 2000）。本文选择具有代表性的 36 颗锆石进行了 LA-ICP-MS 测年分析，分析结果显示测点数据谐和度高且较为集中，均落在谐和线上及其附近。 $^{206}\text{Pb}/^{238}\text{U}$  年龄介于  $275 \pm 4 \sim 269 \pm 3$  Ma 之间，加权平均年龄为  $271 \pm 3$  Ma，MSWD = 0.027（图 7b），时代为中二叠世早期。

表2 辉长岩锆石LA-ICP-MS U-Pb同位素测年结果  
Table 2 LA-ICP-MS zircon U-Pb isotopic dating results of gabbro

样点编号	$w_B/10^{-6}$		比值						年龄/Ma			谐和度								
	Th/U		$^{207}\text{Pb}/^{206}\text{Pb}$		$^{207}\text{Pb}/^{235}\text{U}$		$^{206}\text{Pb}/^{238}\text{U}$		$^{207}\text{Pb}/^{232}\text{Th}$											
	Th	U	$1\sigma$	$207\text{Pb}/^{206}\text{Pb}$	$1\sigma$	$^{207}\text{Pb}/^{235}\text{U}$	$1\sigma$	$^{206}\text{Pb}/^{238}\text{U}$	$1\sigma$	$^{207}\text{Pb}/^{232}\text{Th}$	$1\sigma$									
1	148	144	1.03	0.05372	0.00179	0.31616	0.01011	0.04268	0.00050	0.01038	0.00051	359	73	279	8	269	3	209	10	1.04
2	499	645	0.77	0.05130	0.00082	0.30297	0.00437	0.04283	0.00039	0.00978	0.00032	254	36	269	3	270	2	197	6	0.99
3	450	320	1.41	0.05258	0.00116	0.31135	0.00641	0.04294	0.00043	0.01053	0.00036	311	49	275	5	271	3	212	7	1.02
4	434	309	1.40	0.04959	0.00112	0.29425	0.00621	0.04302	0.00043	0.01025	0.00036	176	52	262	5	272	3	206	7	0.96
5	140	109	1.29	0.05065	0.00284	0.30176	0.01639	0.04320	0.00069	0.00955	0.00067	225	124	268	13	273	4	192	13	0.98
6	202	211	0.95	0.04787	0.00141	0.28678	0.00810	0.04344	0.00047	0.00960	0.00045	92	70	256	6	274	3	193	9	0.93
7	189	285	0.66	0.05009	0.00149	0.29887	0.00841	0.04327	0.00050	0.00985	0.00063	199	68	266	7	273	3	198	13	0.97
8	42	44	0.94	0.05140	0.00524	0.30380	0.03027	0.04286	0.00103	0.00973	0.00147	259	218	269	24	271	6	196	29	1.00
9	216	236	0.92	0.05246	0.00147	0.31116	0.00827	0.04302	0.00048	0.01035	0.00052	305	63	275	6	272	3	208	10	1.01
10	156	166	0.94	0.05366	0.00179	0.31762	0.01012	0.04293	0.00053	0.01098	0.00061	357	74	280	8	271	3	221	12	1.03
11	230	298	0.77	0.04992	0.00128	0.29683	0.00718	0.04312	0.00046	0.01070	0.00053	191	59	264	6	272	3	215	11	0.97
12	260	357	0.73	0.05123	0.00114	0.30510	0.00636	0.04319	0.00043	0.01039	0.00049	251	50	270	5	273	3	209	10	0.99
13	206	197	1.05	0.05274	0.00161	0.31431	0.00910	0.04322	0.00050	0.01042	0.00057	318	68	278	7	273	3	210	11	1.02
14	147	140	1.05	0.04936	0.00191	0.29279	0.01095	0.04302	0.00054	0.01085	0.00065	165	88	261	9	272	3	218	13	0.96
15	260	263	0.99	0.05216	0.00145	0.31016	0.00814	0.04313	0.00048	0.01122	0.00059	292	62	274	6	272	3	226	12	1.01
16	38	74	0.52	0.05172	0.00290	0.30538	0.01665	0.04282	0.00065	0.01021	0.00102	273	124	271	13	270	4	205	20	1.00
17	74	102	0.73	0.05482	0.00241	0.32357	0.01376	0.04281	0.00057	0.01101	0.00089	405	95	285	11	270	4	221	18	1.05
18	62	99	0.62	0.04968	0.00242	0.29807	0.01402	0.04351	0.00064	0.01352	0.00117	180	110	265	11	275	4	272	23	0.96
19	42	71	0.59	0.04746	0.00289	0.28398	0.01692	0.04340	0.00065	0.01160	0.00122	72	140	254	13	274	4	233	24	0.93
20	84	142	0.59	0.05242	0.00184	0.31212	0.01049	0.04318	0.00052	0.01112	0.00080	304	78	276	8	273	3	224	16	1.01
21	647	560	1.15	0.05098	0.00092	0.30435	0.00502	0.04330	0.00040	0.01212	0.00050	240	41	270	4	273	2	244	10	0.99
22	231	317	0.73	0.05218	0.00134	0.30949	0.00745	0.04302	0.00046	0.01318	0.00074	293	57	274	6	272	3	265	15	1.01
23	828	589	1.41	0.04945	0.00088	0.29263	0.00476	0.04293	0.00039	0.01255	0.00052	169	41	261	4	274	4	233	24	0.93
24	563	440	1.28	0.05180	0.00103	0.30853	0.00564	0.04321	0.00042	0.01203	0.00055	277	45	273	4	273	3	242	11	1.00
25	501	664	0.75	0.04968	0.00083	0.29502	0.00451	0.04308	0.00039	0.01265	0.00061	180	39	263	4	272	2	254	12	0.97
26	551	414	1.33	0.05114	0.00104	0.30280	0.00569	0.04296	0.00041	0.01298	0.00063	247	46	269	4	271	3	261	13	0.99
27	658	537	1.23	0.05219	0.00094	0.30873	0.00510	0.04292	0.00040	0.01274	0.00062	294	41	273	4	271	2	252	12	0.96
28	201	180	1.12	0.05144	0.00173	0.30372	0.00981	0.04283	0.00050	0.01379	0.00089	261	76	269	8	270	3	277	18	1.00
29	123	165	0.75	0.05301	0.00178	0.31301	0.00998	0.04284	0.00053	0.01494	0.00110	329	74	277	8	270	3	300	22	1.02
30	438	352	1.24	0.04873	0.00112	0.28767	0.00621	0.04283	0.00042	0.01312	0.00072	135	53	257	5	270	3	264	14	0.95
31	742	442	1.68	0.05082	0.00099	0.30066	0.00543	0.04292	0.00040	0.01371	0.00072	233	45	267	4	271	2	275	14	0.99
32	631	605	1.04	0.05049	0.00090	0.29795	0.00486	0.04281	0.00039	0.01381	0.00076	218	41	265	4	270	2	277	15	0.98
33	289	363	0.79	0.05062	0.00122	0.29711	0.00671	0.04259	0.00044	0.01384	0.00089	223	55	264	5	269	3	278	18	0.98
34	127	266	0.48	0.05058	0.00129	0.29778	0.00715	0.04271	0.00044	0.01467	0.00110	222	58	265	6	270	3	294	22	0.98
35	43	76	0.57	0.04639	0.00318	0.27578	0.01854	0.04313	0.00066	0.01235	0.00163	18	157	247	15	272	4	248	33	0.91
36	239	262	0.91	0.04822	0.00127	0.28538	0.00711	0.04294	0.00044	0.01486	0.00097	110	61	255	6	271	3	298	19	0.94

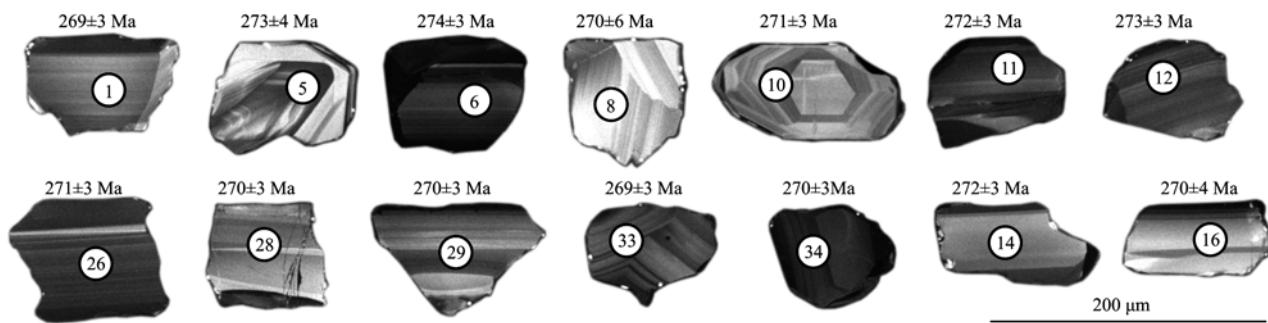


图6 赛什腾辉长岩典型锆石阴极发光图像

Fig. 6 CL images of the representative zircons for gabbro of Saishiteng Mountain

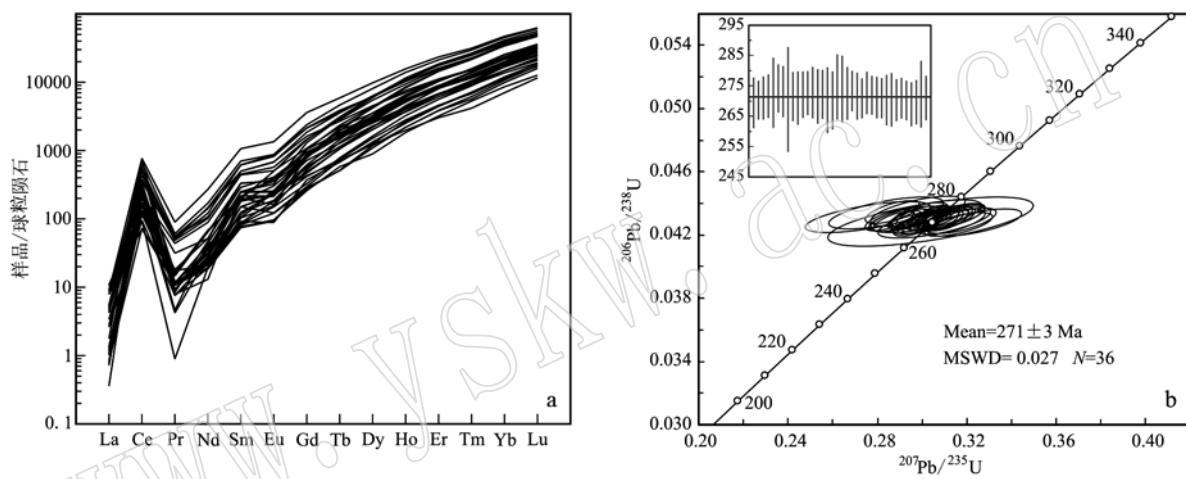


图7 辉长岩锆石稀土元素球粒陨石标准化配分图(a)及锆石U-Pb年龄谐和图(b)

Fig. 7 The chondrite-normalized REE patterns (a) and the U-Pb concordian diagram (b) of zircons for gabbro

## 5 讨论

由于蚀变作用和变质作用通常会导致大离子亲石元素(如K、Rb、Sr、Ba、Cs、 $Pb^{2+}$ 、 $Eu^{2+}$ )具有明显的活动性,而稀土元素以及部分高场强元素(Nb、Ta、Zr、Hf、Th、REE、Ce、U、Ti)甚至在高级变质作用中亦能相对稳定(Hajash, 1984; Becker *et al.*, 1999; Escuder-Viruete *et al.*, 2010),故本文主要利用不活动元素来进行相关讨论。

### 5.1 岩石成因

#### 5.1.1 同化混染与分离结晶

赛什腾辉长岩与围岩侵入关系明显,露头未见同化混染,在锆石CL图像及年龄上也未发现继承或捕获锆石,表明岩浆侵位过程中未遭受或有弱的同化混染作用。原始地幔标准化Th/Nb值( $\gg 1$ )

(Ormerod *et al.*, 1988)和Nb/La值( $< 1$ )(Ernst *et al.*, 2000)是辨别地壳混染作用的两个可靠的微量元素指标。赛什腾辉长岩( $Th/Nb$ )<sub>PM</sub>值为3.26~5.05,Nb/La值为0.39~0.58,在Nb/La-(Th/Nb)<sub>PM</sub>图解中落入遭受地壳混染范围内,暗示辉长岩形成过程中遭受一定程度的地壳混染。Neal等(2002)提出可以用( $La/Nb$ )<sub>PM</sub>和( $Th/Ta$ )<sub>PM</sub>值来区分上地壳和下地壳物质的混染作用,图8b显示样品落至平均上地壳与平均下地壳之间并偏向上地壳一侧,暗示辉长岩岩浆上升过程中可能遭受了中上地壳物质的混染。此外,总分配系数相同或相近的元素的比值受部分熔融和分离结晶作用影响较小,因此可依据总分配系数相同或相近、对地壳混染作用又敏感的元素比值(如Ce/Pb、Ta/Yb、La/Nb、Th/Ta、Nb/Ta、Ti/Yb、Zr/Nb、La/Yb)之间的协变关系,来判断是否存在地壳混染作用以及混染程度高低(Camp-

bell and Griffiths, 1993; Baker *et al.*, 1997; McDonald *et al.*, 2001; 夏明哲等, 2010)。由图9可知, 赛什腾辉长岩除  $\text{Ta}/\text{Yb} - \text{Ce}/\text{Pb}$  ( $R = 0.95$ ) 相关性较好外,  $\text{Th}/\text{Ta} - \text{La}/\text{Nb}$  ( $R = 0.46$ )、 $\text{Ti}/\text{Yb} - \text{Nb}/\text{Ta}$  ( $R = 0.09$ )、 $\text{La}/\text{Yb} - \text{Zr}/\text{Nb}$  ( $R = 0.19$ ) 的相关性均较

差, 表明岩浆演化过程中遭受了地壳混染, 但混染程度整体较弱。

分离结晶作用是镁铁质岩浆分异演化最重要的机制。赛什腾辉长岩  $\text{Mg}^{\#}$  为  $47.82 \sim 61.82$ , 小于原始岩浆玄武岩的  $\text{Mg}^{\#}$  值 ( $68 \sim 78$ ),  $\text{Cr}$  ( $35.9 \sim 105$ )

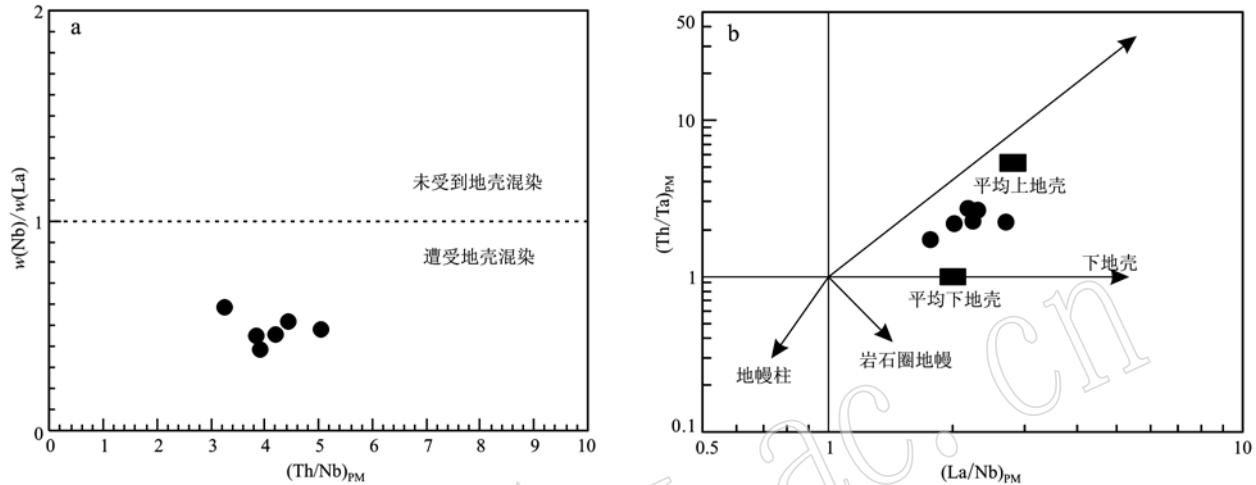


图8 赛什腾辉长岩  $\text{Nb}/\text{La} - (\text{Th}/\text{Nb})_{\text{PM}}$  图解(a, 底图据夏林圻等, 2007)和  $(\text{Th}/\text{Ta})_{\text{PM}} - (\text{La}/\text{Nb})_{\text{PM}}$  图解(b, 底图据 Neal *et al.*, 2002)

Fig. 8  $\text{Nb}/\text{La}$  versus  $(\text{Th}/\text{Nb})_{\text{PM}}$  diagram (a, after Xia Linqi *et al.*, 2007) and  $(\text{Th}/\text{Ta})_{\text{PM}}$  versus  $(\text{La}/\text{Nb})_{\text{PM}}$  diagram (b, after Neal *et al.*, 2002) for gabbro of Saishiteng Mountain

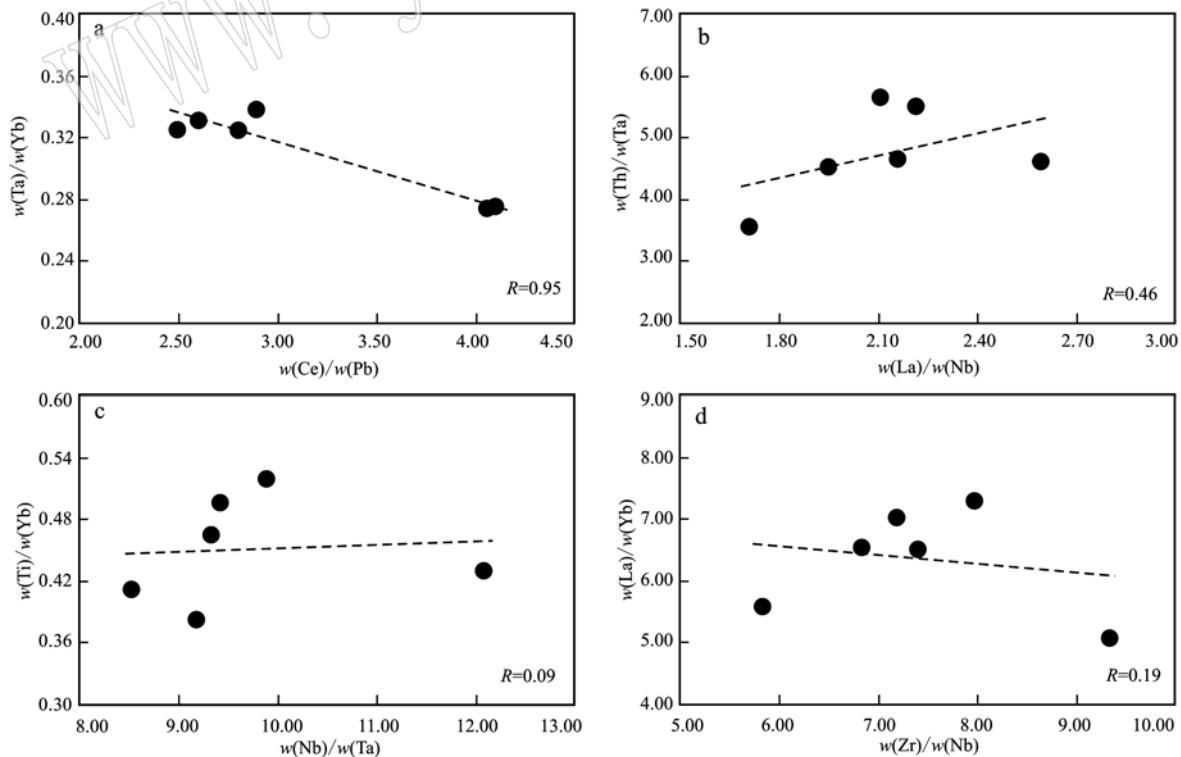


图9 辉长岩微量元素同化混染作用程度判别图

Fig. 9 Discrimination plots for contamination from selected trace elements of gabbro

$<300$  和 Ni(24.9~78.8)  $<250$ , 除一个样品(8.45%)外, 其余 MgO 含量均小于 8%, 均暗示其岩浆在演化过程中经历了一定程度的分离结晶作用(Righter, 2000)。在哈克图解中, MgO 与  $\text{SiO}_2$ 、 $\text{Na}_2\text{O} + \text{K}_2\text{O}$  以及  $\text{Al}_2\text{O}_3$  存在负相关关系(图 3a、3b、3d), 显示辉长岩原始岩浆可能存在斜长石与单斜辉石的分离结晶, 而辉长岩弱(或无)的  $\delta\text{Eu}$  负异常(0.85~1.00)则表明斜长石分离结晶的程度较弱。Spath 等(2001)认为在有斜长石存在的条件下, 若有相当数量单斜辉石的分离结晶作用, 则  $\text{CaO}/\text{Al}_2\text{O}_3$  会随着  $\text{Mg}^{\#}$  比值的增加而增加。辉长岩的  $\text{Mg}^{\#}$  与  $\text{CaO}/\text{Al}_2\text{O}_3$  存在明显的正相关性(图 3f), 表明岩浆结晶分异过程中单斜辉石是主要的结晶分离产物, 这与辉长岩中存在大量粗粒单斜辉石是一致的。

### 5.1.2 岩浆源区性质

因不相容元素具有相似的分配系数, 受岩浆分离结晶和地幔部分熔融作用影响较小, 故通常用来分析源区特征(Taylor and McLennan, 1985)。Shaw 等(2003)认为石榴石橄榄岩源区和尖晶石橄榄岩源区具有不同的稀土元素矿物相/熔体相分配系数, 不管是石榴石橄榄岩还是尖晶石橄榄岩的部分熔融都会使熔体中的轻稀土元素富集, 且随着熔融程度的不同, 源区轻、重稀土元素比值也会发生变化, 并因重稀土元素更容易进入石榴石中而导致这种变化在石榴石橄榄岩中更大一些, 故可以用稀土元素比值(或原始地幔标准化)来限定地幔岩浆源区的组分及

部分熔融的程度。赛什腾辉长岩低的  $(\text{La/Yb})_{\text{N}}$  值(3.44~4.94)、 $(\text{Gd/Yb})_{\text{N}}$  值(1.60~1.86)以及 HREE 相对平坦的球粒陨石标准化稀土元素配分曲线, 暗示辉长岩源区物质可能不含石榴石(蓝江波等, 2007; Pollock and Hibbard, 2010)。Wang 等(2002)认为石榴石橄榄岩平衡熔体具有较高的原始地幔标准化的  $(\text{Tb/Yb})_{\text{PM}}$ ( $>1.8$ ), 而尖晶石橄榄岩则具有低的  $(\text{Tb/Yb})_{\text{PM}}$  值( $<1.8$ )。源区赛什腾辉长岩的  $(\text{Tb/Yb})_{\text{PM}}$  值为 1.50~1.67, 在  $(\text{Tb/Yb})_{\text{PM}}-(\text{La/Sm})_{\text{PM}}$  图解中(图 10a)落入尖晶石橄榄岩区域, 表明岩石部分熔融应发生在尖晶石稳定区域(Wang et al., 2002; 解超明等, 2019)。此外, Deniel(1998)提出 Ce/Y-Zr/Nb 图解可作为判别玄武岩浆源区矿物相组合的有效图解, 在 Ce/Y-Zr/Nb 图解中(图 10b), 样品均位于原始尖晶石相二辉橄榄岩和亏损尖晶石相二辉橄榄岩熔融源区之间, 进一步表明赛什腾辉长岩的岩浆源区为尖晶石二辉橄榄岩, 并暗示其源区深度可能在 70 km 左右(Lambert and Wyllie, 1968)。

综上可知, 赛什腾辉长岩原岩岩浆是地幔尖晶石二辉橄榄岩部分熔融的产物, 在演化过程中经历了弱的斜长石以及较为明显的单斜辉石的分离结晶作用, 并在上升侵位过程中遭受了弱的中上地壳的同化混染作用。

### 5.2 构造环境及地质意义

前已提及, 赛什腾辉长岩具明显富集 Rb、Ba、Th、U、Sr 等大离子亲石元素、显著亏损 Nb、Ta、Zr 高

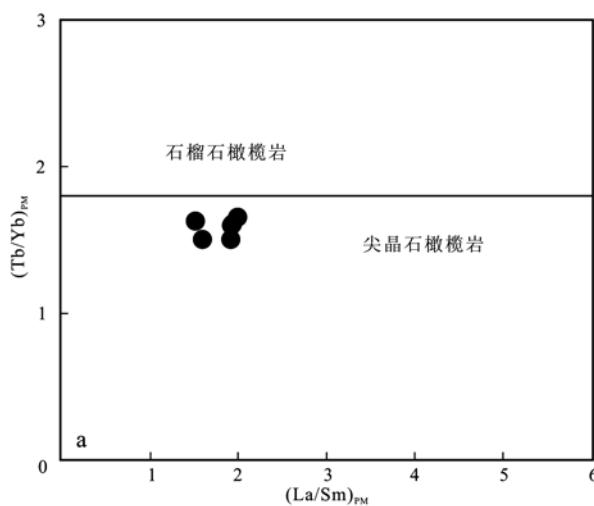
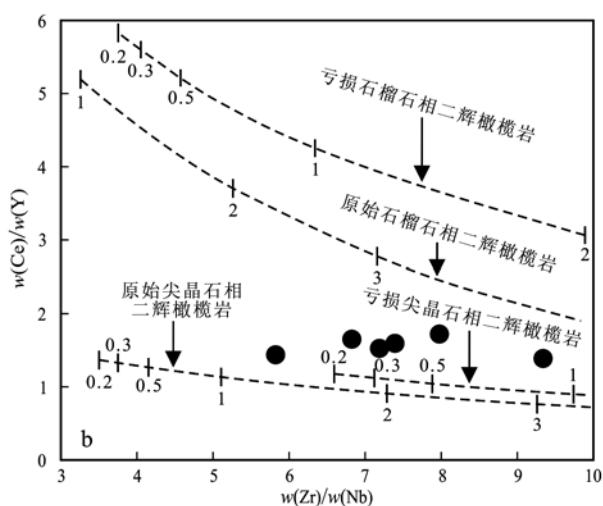


图 10 赛什腾辉长岩  $(\text{Tb/Yb})_{\text{PM}}-(\text{La/Sm})_{\text{PM}}$  图解(a, 底图据 Wang et al., 2002)和 Ce/Y-Zr/Nb 图解(b, 底图据 Deniel, 1998)

Fig. 10  $(\text{Tb/Yb})_{\text{PM}}$  versus  $(\text{La/Sm})_{\text{PM}}$  diagram (a, after Wang et al., 2002) and Ce/Y versus Zr/Nb diagram (b, after Deniel, 1998) for gabbro of Saishiteng Mountain



场强元素以及轻稀土元素相对富集、重稀土元素平坦且向右缓倾的稀土元素配分模式的地球化学特征,具有俯冲消减带弧火山岩的特征,显示其成因可能与消减作用有关(欧阳京等,2010)。而正常的岛弧或活动陆缘弧玄武岩以及受到地壳或岩石圈混染的软流圈(或地幔柱)源大陆玄武岩均可出现上述似消减带信号的地球化学特征。夏林圻等(2007)认为地壳混染作用对于Zr、Y等元素的原始浓度不会产生重大影响,且岛弧玄武岩总体上是以具有较低的Zr含量( $<130 \times 10^{-6}$ )和Zr/Y值( $<4$ )为特征,大陆玄武岩不管是否遭受地壳或岩石圈混染都具有较高的Zr含量( $>70 \times 10^{-6}$ )和Zr/Y值( $>3$ )。样品中Zr含量为 $32.9 \times 10^{-6} \sim 59.1 \times 10^{-6}$ , Zr/Y=1.68~2.95,符合岛弧玄武岩低Zr含量( $<130 \times 10^{-6}$ )

和Zr/Y( $<4$ )的特征,在Zr/Y-Zr图解中(图11a),样品也均落入火山弧玄武岩区域,表明赛什腾辉长岩应形成于与消减作用有关的岛弧构造环境。进一步利用Th/Yb-Ta/Yb图解(Pearce, 1982)对弧型岩浆岩进行细分(图11b),辉长岩落入活动大陆边缘玄武岩(VAB)类型范围内,这与岩石中相对较高的Th/Nb值(0.22~0.67)和Th×Ta/Zr<sup>2</sup>值(0.0005~0.0018)一致(前者 $>0.1$ ,后者 $>0.0005$ ;李永军等,2015),都指示辉长岩具有陆缘弧玄武岩的特征。此外,辉长岩富MgO及FeO<sup>T</sup>、贫Alk和P<sub>2</sub>O<sub>5</sub>、高Al<sub>2</sub>O<sub>3</sub>含量(17.31%~19.45%)的主要元素特征,与典型高铝玄武岩相似(桑隆康等,2012;李永军等,2019),也同样表明辉长岩形成于活动大陆边缘岛弧环境。

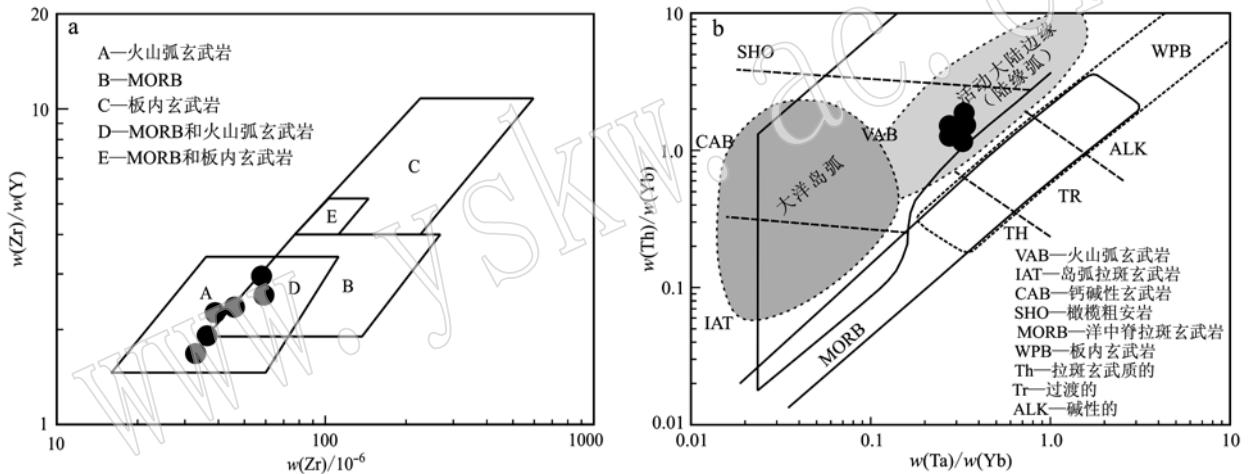


图11 赛什腾辉长岩Zr/Y-Zr图解和Th/Yb-Ta/Yb图解(底图据 Pearce, 1982)

Fig. 11 Zr/Y versus Zr and Th/Yb versus Ta/Yb diagrams for gabbro of Saishiteng Mountain (after Pearce, 1982)

柴北缘构造带由南向北可划分为柴北缘早古生代俯冲带、欧龙布鲁克地块以及宗务隆构造带3个次一级构造单元,其中宗务隆构造带被认为是一独立演化发展的印支期造山带,它经历了由早泥盆世的陆内裂陷(孙延贵等,2004)、晚石炭世的洋盆发育(王毅智等,2001)和晚二叠世到中三叠世的俯冲-碰撞造山的演化过程(强娟,2008;郭安林等,2009;彭渊等,2016)。近年来,在冷湖北山、小赛什腾山以及赛什腾山三岔沟一带发现了一系列具岛弧或活动大陆边缘弧性质的中二叠世花岗岩,且均被认为是宗务隆洋向南俯冲的产物(吴锁平,2008;邱士东等,2015;高万里等,2019)。赛什腾辉长岩位于柴北缘构造带欧龙布鲁克地块西北部,LA-ICP-MS锆石U-Pb测年显示其形成于中二叠世早期( $271 \pm 3$

Ma),与上述花岗岩相似,形成于活动大陆边缘岛弧环境。结合晚古生代晚期区域地质演化特征,认为赛什腾辉长岩可能为中二叠世早期宗务隆洋壳向欧龙布鲁克地块南向俯冲的产物,表明宗务隆有限洋盆向西延伸至冷湖北山-赛什腾山一带,并在中二叠世早期之前已经发生洋壳俯冲作用。然而,红柳沟北-八罗根郭勒河一带中二叠世( $266.8 \pm 1.5$  Ma)的辉绿玢岩-辉长岩-枕状玄武岩-细碧岩-硅质岩组合被认为是宗务隆构造带当时洋壳物质的一部分,代表着中二叠世宗务隆构造带中部仍处于洋盆发育阶段(许海全等,2019)。天峻南山及其以东地区广泛出露的晚二叠世-中三叠世的活动大陆边缘岛弧型花岗岩(强娟,2008;郭安林等,2009;彭渊等,2016)则暗示宗务隆构造带东段晚二叠世-中三叠世仍处

于洋壳俯冲演化阶段,而同时期小赛什腾山S型花岗岩的形成则表明该地区中三叠世早期( $242.6 \pm 3.2$  Ma)已经发展俯冲结束后的同碰撞阶段(杨明慧等,2002)。通过上述分析可知,宗务隆构造带西段俯冲-碰撞事件明显早于东段,而东西两段俯冲-碰撞时限的差异表明宗务隆洋盆存在西早东晚的“剪刀式闭合”的可能。

## 6 结论

(1) 锆石LA-ICP-MS U-Pb测年结果显示赛什腾辉长岩的结晶年龄为 $271 \pm 3$  Ma(MSWD = 0.027),表明辉长岩形成时代为中二叠世早期。

(2) 岩石地球化学特征显示,赛什腾辉长岩源于地幔尖晶石二辉橄榄岩的部分熔融,岩浆在演化过程经历了一定程度的斜长石和单斜辉石结晶分异作用,并遭受了弱的中上地壳物质的同化混染。

(3) 赛什腾辉长岩形成于活动大陆边缘岛弧环境,是中二叠世早期宗务隆洋西段南向俯冲的产物。结合区域构造演化可知,晚古生代晚期宗务隆构造带俯冲碰撞事件存在西早东晚的演化特征。

## References

- Andersen T. 2002. Correction of common lead in U-Pb analyses that do not report  $^{204}\text{Pb}$ [J]. Chemical Geology, 192: 59~79.
- Baker J A, Menzies M A, Thirlwall M F, et al. 1997. Petrogenesis of Quaternary intraplate volcanism, Sana'a Yenmen: Implication and polybaric melt hybridization[J]. Journal of Petrology, 38: 1359~2390.
- Becker H, Jochum K P and Carlson R W. 1999. Constraints from high-pressure veins in eclogites on the composition of hydrous fluids in subduction zones[J]. Chemical Geology, 160(4): 291~308.
- Campbell I H and Griffiths R W. 1993. The evolution of mantle's chemical structure[J]. Lithos, 30: 389~399.
- Chen Danling, Sun Yong and Liu Liang. 2007. The metamorphic ages of the country rock of the Yukahe eclogites in the North Qaidam and its geological significance[J]. Earth Science Frontiers, 14(1): 108~116(in Chinese with English abstract).
- Cox K J, Bell J D and Pankhurst R J. 1979. The Interpretation of Igneous Rocks[M]. Allen and Unwin, St Leonards, NSW, Australia.
- Deniel C. 1998. Geochemical and isotopic(Sr, Nd, Pb) evidence for plume-lithosphere interactions in the genesis of Grande Comore magmas (Indian Ocean)[J]. Chemical Geology, 144(3~4): 281~303.
- Dong Zengchan, Gu Pingyang, Chen Ruiming, et al. 2015b. Geochronology, geochemistry and Hf isotope of Yanchangbeishan Adamellite of Lenghu Area in Qinghai[J]. Earth Science Journal of China University of Geosciences, 40(1): 130~144(in Chinese with English abstract).
- Dong Zengchan, Gu Pingyang, Jiao He, et al. 2014b. Geochemistry and chronology of Yanchang beishan gabro in Lenghu area at the west segment of the north margin of Qaidam[J]. Chinese Journal of Geology, 49(4): 1132~1149(in Chinese with English abstract).
- Dong Zengchan, Xiao Peixi, Gu Pingyang, et al. 2014a. Geochronology, geochemistry and Hf isotope characteristics of yanchang beishan granodiorite of lenghu area in Qinghai[J]. Northwestern Geology, 47(4): 141~155(in Chinese with English abstract).
- Dong Zengchan, Yang Cheng, Gu Pingyang, et al. 2015a. Geochronology and geochemistry of Yanchang beishan biotite Adamellite of Lenghu area in Qinghai and their geological significance[J]. Geotectonica et Metallogenesis, 39(1): 167~178(in Chinese with English abstract).
- Ernst R E, Buchan K L, Hamilton M A, et al. 2000. Integrated paleomagnetism and U-Pb geochronology of mafic dikes of the eastern Anabar Shield Region, Siberia: Implications for Mesoproterozoic Paleolatitude of siberia and comparison with Laurentia[J]. The Journal of Geology, 108: 381.
- Escuder-Viruete J, Pérez-Estaún A, Weis D, et al. 2010. Geochemical characteristics of the Rio Verde complex, central Hispaniola: Implications for the paleotectonic reconstruction of the Lower Cretaceous Caribbean island-arc[J]. Lithos, 114(1~2): 168~185.
- Gao Wanli, Wang Zongxiu, Li Leilei, et al. 2019. Discovery of the Permian granite in saishiteng mountain of the northern qaidam basin and its tectonic significance[J]. Acta Geologica Sinica, 93(4): 816~829(in Chinese with English abstract).
- Gao Xiaofeng, Xiao Peixi and Jia Qunzi, 2011. Redetermination of the Tanjianshan Group: Geochronological and geochemical evidence of Basalts from the margin of the Qaidam Basin[J]. Acta Geologica Sinica, 85(9): 1452~1463(in Chinese with English abstract).
- Gu Pingyang, Chen Ruiming, Zha Xianfeng, et al. 2018. The age, petrogenesis and geological significance of quartz diorite on the northwestern margin of Qaidam basin[J]. Acta Petrologica et Mineralogica, 37(1): 19~33(in Chinese with English abstract).
- Guo Anlin, Zhang Guowei, Qiang Juan, et al. 2009. Indosinian Zongwu-

- long orogenic belt on the northeastern margin of the Qinghai-Tibet plateau[J]. *Acta Petrologica Sinica*, 25(1): 1~12 (in Chinese with English abstract).
- Hajash A J. 1984. Rare earth element abundances and distribution patterns in hydrothermally altered basalts: Experimental results[J]. *Contributions to Mineralogy and Petrology*, 85(4): 409~412.
- Hoskin P W O and Black L P. 2000. Metamorphic zircon formation by solid-state recrystallization of protolith igneous zircons[J]. *J. Meta. Geol.*, 18: 423~439.
- Lambert I B and Wyllie P J. 1968. Stability of hornblende and a model for the low velocity zone[J]. *Nature*, 219(5 160): 1 240~1 241.
- Lan Jiangbo, Xu Yigang, Yang Qijun, et al. 2007. ~40 Ma OIB-type mafic magmatism in the Gaoligong belt: results of break-off between subducting Tethyan slab and Indian plate? [J]. *Acta Petrologica Sinica*, 23(6): 1 334~1 346 (in Chinese with English abstract).
- Li Yongjun, Li Ganyu, Tong Lili, et al. 2015. Discrimination of ratios of Ta, Hf, Th, La, Zr and Nb for tectonic settings in basalts[J]. *Journal of Earth Sciences and Environment*, 37(3): 14~21 (in Chinese with English abstract).
- Li Yongjun, Tao Xiaoyang, Wang Zuopeng, et al. 2019. Discovery of high-alumina basalt of Tuoranggekudu Formation in Yematan of Qinghe Area, East Junggar, China and its geological significance [J]. *Journal of Earth Sciences and Environment*, 41(1): 1~14 (in Chinese with English abstract).
- Li Wenxuan, Wu Xinguo and Feng Jialin. 1994. Study on the basalt of volcania cluster in Datong, Shanxi[J]. *Journal of Hebei College of Geology*, 6: 547~555 (in Chinese with English abstract).
- Ludwig K R. 2003. 3. 0—A geochronological toolkit for Micro-soft Excel [J]. *Berkeley Geochronology Center, Special Publication*, 4: 1~70.
- Mcdonald R, Rogers N W, Fitton J G, et al. 2001. Plume-lithosphere interactions in the generation of the basalts of the Kenya Rift, East Africa[J]. *Journal of Petrology*, 42: 877~900.
- Meng Fancong, Zhang Jianxin, Yang Jingsui, et al. 2003. Geochemical characteristics of eclogites in Xitieshan area, North Qaidam of northwestern China[J]. *Acta Geologica Sinica*, 19(3): 435~442 (in Chinese with English abstract).
- Miyashiro A. 1974. Volcanic rock series in island arcs and active continental margins[J]. *American Journal of Science*, 274(4): 321~355.
- Neal C R, Mahoney J J and Chazey W J. 2002. Mantle sources and the highly variable role of continental lithosphere in basalt petrogenesis of the Kerguelen Plateau and Broken Ridge LIP: Results from ODP Leg 183[J]. *Journal of Petrology*, 43: 1 177~1 205.
- Ormerod D S, Hawkesworth C J, Rogers N W, et al. 1988. Tectonic and magmatic transition in the western Great Basin, USA [J]. *Nature*, 333(6 171~6 172): 394~353.
- Quyang Jing, Wang Shuangshuang, Yu Man, et al. 2010. Igneous assemblages of different origin in island arc and their geological significances[J]. *Gansu Geology*, 19(2): 18~26 (in Chinese with English abstract).
- Pan Guitang, Li Xingzhen, Wang Liquan, et al. 2002. Preliminary division of tectonic units of the Qinghai-Tibet Plateau and its adjacent regions[J]. *Geological Bulletin of China*, 21(11): 701~707 (in Chinese with English abstract).
- Pearce J A. 1982. Trace element characteristics of lavas from destructive plate boundaries[A]. Thorpe R S. *Andesites*[C]. Chichester: Wiley, 525~548.
- Peng Yuan, Ma Yansheng, Liu Chenglin, et al. 2016. Geological characteristics and tectonic significance of the Indosinian granodiorites from the Zongwulong tectonic belt in North Qaidam[J]. *Earth Science Frontiers*, 23(2): 206~221 (in Chinese with English abstract).
- Pollock J C and Hibbard J P. 2010. Geochemistry and tectonic significance of the Stony Mountain gabbro, North Carolina: Implications for the Early Paleozoic evolution of Carolinia[J]. *Gondwana Research*, 17(2~3): 500~515.
- Qiang Juan. 2008. The Granitoids in Zongwulong Tectonic Zone on the Northeastern Margin of the Qinghai-Tibet Plateau and Its tectonic Significance[D]. Xi'an: Northwest University (in Chinese with English abstract).
- Qiu Shidong, Gu Pingyang, Pang Xinyu, et al. 2015. Zircon MC-LA-ICP-MS U-Pb dating, Geochemistry and Geological Significance of Gneissic Quartz Diorite in Northern Lenghu, Qinghai[J]. *Geological Review*, 61(4): 948~960 (in Chinese with English abstract).
- Righter K A. 2000. A comparison of basaltic volcanism in the Cascades and western Mexico: Compositional diversity in continental arcs[J]. *Tectonophysics*, 318(1~4): 99~117. Doi: 10. 1016/S0040-1951(99)00308-X.
- Shi Rendeng, Yang Jingsui and Wu Cailai. 2003. The discovery of adakitic dacite in Early Palaeozoic island arc volcanic rocks on the northern margin of Qaidam basin and its geological significance[J]. *Acta Petrologica et Mineralogica*, 22(3): 229~236 (in Chinese with English abstract).
- Sang Longkang and Ma Changqian. 2012. *Petrology*[M]. Beijing: Geological Publishing House, 1~620 (in Chinese).
- Shaw J E, Baker J A, Menzies M A, et al. 2003. Petrogenesis of the largest intraplate volcanic field on the Arabian Plate (Jordan): A mixed

- lithosphere asthenosphere source activated by lithospheric extension [J]. *Journal of Petrology*, 44(9): 1 657 ~ 1 679.
- Song Shuguang, Wang Mengjue, Wang Chao, et al. 2015. Magmatism during continental collision, subduction, exhumation and mountain collapse in collisional orogenic belts and continental net growth: A perspective[J]. *Science China: Earth Sciences*, 58: 1 284 ~ 1 304 (in Chinese with English abstract).
- Song Shuguang and Yang Jingsui. 2001. Sanidine + quartz inclusions in Dulan eclogites: Evidence for UHP metamorphism on the North margin of the Qaidam Basin, N W China[J]. *Acta Geologica Sinica*, 75(2): 180 ~ 185 (in Chinese with English abstract).
- Spath A, Roex A P and Opiyo A N. 2001. Plume-lithosphere interaction and the origin of continental rift-related alkaline volcanism—the Chyulu Hills Volcanic Province, Southern Kenya[J]. *Journal of Petrology*, 42: 765 ~ 778.
- Sun S S and McDonough W F. 1989. Chemical and isotopic systematics of oceanic basalts: implications for mantle composition and process [A]. Saunders A D and Norry M J. *Magmatism in the Ocean Basins* [C]. Geological Society Special Publication, 42: 313 ~ 345.
- Sun Yangui, Zhang Guowei, Wang Jin, et al. 2004.  $^{40}\text{Ar}/^{39}\text{Ar}$  age of the basic sill swarms of two periods in the junction area of Qinling and Kunlun and its tectonic significance[J]. *Acta Geologica Sinica*, 78(1): 65 ~ 71 (in Chinese with English abstract).
- Taylor S R and McLennan S M. 1985. The continental crust: Its composition and evolution[J]. Oxford: Blackwell Scientific Publications, 1 ~ 312.
- Wang Huichu, Lu Songnian, Mo Xuanxue, et al. 2005. An Early Paleozoic collisional orogen on the northern margin of the Qaidam basin, northwestern China[J]. *Geological Bulletin of China*, 24(7): 603 ~ 612 (in Chinese with English abstract).
- Wang K, Plant T, Walker J D, et al. 2002. A mantle melting profile across the Basin and Range, SW USA[J]. *Journal of Geophysical Research: Solid Earth*, 107(B1): 1 ~ 21.
- Wang Suli and Zhou Lifan. 2016. LA-ICP-MS zircon U-Pb dating, geochemistry and tectonic implication of the bojite in the Zongwulong Mountain[J]. *Journal of Northwest University(Natural Science Edition)*, 46(5): 716 ~ 724 (in Chinese with English abstract).
- Wang Yizhi, Bai Yongshan and Lu Hailian. 2001. Geological characteristics of Tianjunnanshan ophiolite in Qinghai and its forming environment[J]. *Qinghai Geology*, 10(1): 29 ~ 35.
- Wilson M. 1989. Igneous Petrogenesis A Global Tectonic Approach[M]. Kluwer Academic Publishers, London: Unwin Hyman, 1 ~ 466.
- Wu Cailai, Gao Yuanhong, Wu Suoping, et al. 2007. Zircon SHRIMP U-Pb dating of granites from the Da Qaidam area in the north margin of Qaidam basin, NW China[J]. *Acta Petrologica Sinica*, 23(8): 1 861 ~ 1 875 (in Chinese with English abstract).
- Wu Cailai, Gao Yuanhong, Wu Suoping, et al. 2008. Zircon SHRIMP U-Pb dating and geochemical characteristics of granites from the western part of the north margin of Qaidam basin[J]. *Acta Petrologica Sinica*, 38(8): 930 ~ 949 (in Chinese with English abstract).
- Wu Suoping. 2008. The Petrogenesis of Paleozoic Granitoids in the North Margin of Qaidam Basin and their Orogenic Response[D]. Beijing: Chinese Academy of Geological Sciences (in Chinese with English abstract).
- Xia Linqi, Xia Zuchun, Xu Xueyi, et al. 2007. The discrimination between continental basalt and island arc basalt based on geochemical method[J]. *Acta Mineralogica et Petrologica*, 26(1): 77 ~ 89 (in Chinese with English abstract).
- Xia Mingzhe, Jiang Changyi, Qian Zhuangzhi, et al. 2010. Geochemistry and petrogenesis of Huangshandong intrusion, East Tianshan, Xinjiang[J]. *Acta Petrologica Sinica*, 26(8): 2 413 ~ 2 430 (in Chinese with English abstract).
- Xie Chaoming, Duan Menglong, Yu Yunpeng, et al. 2019. Genesis and geological significance of Early Jurassic metamorphic gabbro in the Sumdo area, Tibet[J]. *Acta Petrologica Sinica*, 35(10): 3 065 ~ 3 082.
- Xin Houtian, Wang Huichu and Zhou Shijun. 2006. Geological events and tectonic evolution of the north margin of the Qaidam Basin[J]. *Geological Survey and Research*, 29(4): 311 ~ 320 (in Chinese with English abstract).
- Xu Haiquan, Yan Yufeng, Qu Guangju, et al. 2019. Material composition and evolutionary characteristics of the western section of Zongwulong tectonic belt[J]. *Journal of Qinghai University*, 37(3): 87 ~ 93.
- Xu Xuming, Liu Guang and Liang Guoqing. 2017. LA-ICP-MS zircon U-Pb dating and geochemical characteristics of the volcanic rocks in the area of Saishiteng Mountain[J]. *Xinjiang Geology*, 35(3): 250 ~ 254 (in Chinese with English abstract).
- Yang Jingsui, Song Shuguang, Xu Zhiqin, et al. 2001. Discovery of coesite in the North Qaidam early Paleozoic ultrahigh-high pressure (UHP-HP) metamorphic belt, NW China[J]. *Acta Geologica Sinica*, 75(2): 175 ~ 179 (in Chinese with English abstract).
- Yang Minghui and Song Jianjun. 2002. Petrology of the Lenghu granite mass, northwestern Qaidam basin, China[J]. *Northwestern Geology*, 35(3): 94 ~ 98 (in Chinese with English abstract).
- Zhang C, Zhang L, Herman V R, et al. 2011. Petrology and SHRIMP U-Pb dating of Xiteshan eclogite, North Qaidam UHP metamorphic

- belt, NW China[J]. Journal of Asian Earth Sciences, 42: 752 ~ 767.
- Zhang Yun, Sun Lixin, Zhang Tianfu, et al. 2020. Geochronology, geochemistry and Sr-Nd-Hf isotopic compositions of the Wuhua gabbros in Langshan, Inner Mongolia: Constraints for mantle sources and petrogenesis[J]. Geoscience, 34(3): 450 ~ 465 (in Chinese with English abstract).
- Zhang Z C, Zhou G, Kusky T M, et al. 2009. Late Paleozoic volcanic record of the Northern Junggar Terrane, Xinjiang, northwestern China. Major and trace element characteristics, Sr-Nd isotopic systematica and implications for tectonic evolution[J]. Gondwana Research, 16(2): 201 ~ 215.
- Zhou Bin, Zheng Youye, Xu Rongke, et al. 2013. LA-ICP-MS zircon U-Pb dating and Hf isotope geochemical characteristics of Qaidamshan intrusive body[J]. Geological Bulletin of China, 32(7): 1 027 ~ 1 034 (in Chinese with English abstract).
- Zhuang Yujun, Gu Pingyang, Li Peiqing, et al. 2019. Geochemistry, geochronology and Hf isotopic compositions of metagabbro dykes on the northwestern margin of Oulongbuluke micro-block on the northern margin of Qaidam Basin[J]. Geological Bulletin of China, 38(11): 1 801 ~ 1 812 (in Chinese with English abstract).
- 木盆地周缘玄武岩年代学和地球化学证据[J]. 地质学报, 85(9): 1 452 ~ 1 463.
- 辜平阳, 陈锐明, 查显峰, 等. 2018. 柴达木盆地西北缘石英闪长岩的形成时代、岩石成因及地质意义[J]. 岩石矿物学杂志, 37(1): 19 ~ 33.
- 郭安林, 张国伟, 强娟, 等. 2009. 青藏高原东北缘印支期宗务隆造山带[J]. 岩石学报, 25(1): 1 ~ 12.
- 蓝江波, 徐义刚, 杨启军, 等. 2007. 高黎贡带~40 Ma OIB型基性岩浆活动: 消减特提斯洋片与印度板块断离的产物? [J]. 岩石学报, 23(6): 1 334 ~ 1 346.
- 李文宣, 吴新国, 冯家麟. 1994. 大同火山群玄武岩研究[J]. 河北地质学院学报, 6: 547 ~ 555.
- 李永军, 李甘雨, 佟丽莉, 等. 2015. 玄武岩类形成的大地构造环境 Ta、Hf、Th、La、Zr、Nb 比值对比判别[J]. 地球科学与环境学报, 37(3): 14 ~ 21.
- 李永军, 陶晓杨, 王祚鹏, 等. 2019. 东准噶尔青河地区野马滩一带托让格库都克组高铝玄武岩的发现及其地质意义[J]. 地球科学与环境学报, 41(1): 1 ~ 14.
- 孟繁聪, 张建新, 杨经绥, 等. 2003. 柴北缘锡铁山榴辉岩的地球化学特征[J]. 岩石学报, 19(3): 435 ~ 442.
- 欧阳京, 汪双双, 于漫, 等. 2010. 岛弧环境中不同成因的火成岩组合及其地质意义[J]. 甘肃地质, 19(2): 18 ~ 26.
- 潘桂棠, 李兴振, 王立全, 等. 2002. 青藏高原及邻区大地构造单元初步划分[J]. 地质通报, 21(11): 701 ~ 707.
- 彭渊, 马寅生, 刘成林, 等. 2016. 柴北缘宗务隆构造带印支期花岗岩地质特征及其构造意义[J]. 地学前缘, 23(2): 206 ~ 221.
- 强娟. 2008. 青藏高原东北缘宗务隆构造带花岗岩及其构造意义[D]. 西安: 西北大学.
- 邱士东, 辜平阳, 庞新渝, 等. 2015. 青海冷湖北片麻状石英闪长岩的 MC-LA-ICP-MS 锆石 U-Pb 年龄、地球化学特征及地质意义[J]. 地质论评, 61(4): 948 ~ 960.
- 桑隆康, 马昌前. 2012. 岩石学[M]. 北京: 地质出版社, 1 ~ 620.
- 史仁灯, 杨经绥, 吴才来. 2003. 柴北缘早古生代岛弧火山岩中埃达克质英安岩的发现及其地质意义[J]. 岩石矿物学杂志, 22(3): 229 ~ 236.
- 宋述光, 牛耀龄, 张立飞, 等. 2009. 大陆造山运动: 从大洋俯冲到大陆俯冲、碰撞、折返的时限——以北祁连山、柴北缘为例[J]. 岩石学报, 25(9): 2 067 ~ 2 077.
- 宋述光, 王梦珏, 王潮, 等. 2015. 大陆造山带碰撞-俯冲-折返-垮塌过程的岩浆作用及大陆地壳净生长[J]. 中国科学: 地球科学, 45(7): 916 ~ 940.
- 宋述光, 杨经绥. 2001. 柴达木盆地北缘都兰地区榴辉岩中透长石 +

## 附中文参考文献

- 陈丹玲, 孙勇, 刘良. 2007. 柴北缘鱼卡河榴辉岩围岩的变质时代及其地质意义[J]. 地学前缘, 14(1): 108 ~ 116.
- 董增产, 辜平阳, 陈锐明, 等. 2015b. 柴北缘西端盐场北山二长花岗岩年代学、地球化学及其 Hf 同位素特征[J]. 地球科学(中国地质大学学报), 40(1): 130 ~ 144.
- 董增产, 辜平阳, 焦和, 等. 2014b. 柴北缘西端冷湖盐场北山辉长岩地球化学及年代学研究[J]. 地质科学, 49(4): 1 132 ~ 1 149.
- 董增产, 校培喜, 辜平阳, 等. 2014a. 青海冷湖盐场北山花岗闪长岩年代学、地球化学及其 Hf 同位素特征[J]. 西北地质, 47(4): 141 ~ 155.
- 董增产, 杨成, 辜平阳, 等. 2015a. 青海冷湖盐场北山黑云母二长花岗岩年代学、地球化学及其地质意义[J]. 大地构造与成矿学, 39(1): 167 ~ 178.
- 高万里, 王宗秀, 李磊磊, 等. 2019. 柴达木盆地北缘小赛什腾山二叠纪花岗岩的发现及其构造意义[J]. 地质学报, 93(4): 816 ~ 829.
- 高晓峰, 校培喜, 贾群子. 2011. 滩间山群的重新厘定——来自柴达

- 石英包裹体:超高压变质作用的证据[J]. 地质学报, 75(2): 180~185.
- 孙延贵, 张国伟, 王瑾, 等. 2004. 秦昆结合区两期基性岩墙群<sup>40</sup>Ar/<sup>39</sup>Ar定年及其构造意义[J]. 地质学报, 78(1): 65~71.
- 王惠初, 陆松年, 莫宣学, 等. 2005. 柴达木盆地北缘早古生代碰撞造山系统[J]. 地质通报, 24(7): 603~612.
- 王苏里, 周立发. 2016. 宗务隆山角闪辉长岩LA-ICP-MS锆石U-Pb定年、地球化学特征及其地质意义[J]. 西北大学学报(自然科学版), 46(5): 716~724.
- 王毅智, 拜永山, 陆海莲. 2001. 青海天峻南山蛇绿岩的地质特征及其形成环境[J]. 青海地质, 10(1): 29~35.
- 吴才来, 郁源红, 吴锁平, 等. 2007. 柴北缘大柴旦地区古生代花岗岩锆石SHRIMP定年[J]. 岩石学报, 23(8): 1 861~1 875.
- 吴才来, 郁源红, 吴锁平, 等. 2008. 柴北缘西段花岗岩锆石SHRIMP U-Pb定年及其岩石地球化学特征[J]. 中国科学(D辑: 地球科学), 38(8): 930~949.
- 吴锁平. 2008. 柴北缘古生代花岗岩类成因及其造山响应[D]. 北京: 中国地质科学院.
- 夏林圻, 夏祖春, 徐学义, 等. 2007. 利用地球化学方法判别大陆玄武岩和岛弧玄武岩[J]. 岩石矿物学杂志, 26(1): 77~89.
- 夏明哲, 姜常义, 钱壮志, 等. 2010. 新疆东天山黄山东岩体岩石地球化学特征与岩石成因[J]. 岩石学报, 26(8): 2 413~2 430.
- 解超明, 段梦龙, 于云鹏, 等. 2019. 西藏松多地区早侏罗世变质辉长岩的成因及其构造意义[J]. 岩石学报, 35(10): 3 065~3 082.
- 辛后田, 王惠初, 周世军. 2006. 柴北缘的大地构造演化及其地质事件群[J]. 地质调查与研究, 29(4): 311~320.
- 徐旭明, 刘广, 梁国庆. 2017. 赛什腾山一带火山岩LA-ICP-MS锆石U-Pb年龄及其地球化学特征[J]. 新疆地质, 35(3): 250~254.
- 许海全, 严玉峰, 屈光菊, 等. 2019. 宗务隆构造带西段物质组成及演化特征[J]. 青海大学学报, 37(3): 87~93.
- 杨经绥, 许志琴, 李海兵, 等. 1998. 我国西部柴北缘地区发现榴辉岩[J]. 科学通报, 43(14): 1 544~1 549.
- 杨经绥, 宋述光, 许志琴, 等. 2001. 柴达木盆地北缘早古生代高压-超高压变质带中发现典型超高压矿物——柯石英[J]. 地质学报, 75(2): 175~179.
- 杨明慧, 宋建军. 2002. 柴达木盆地冷湖花岗岩体岩石学初步研究[J]. 西北地质, 35(3): 94~98.
- 张云, 孙立新, 张天福, 等. 2020. 内蒙古狼山地区鸟花辉长岩的年代学、地球化学和Sr-Nd-Hf同位素组成: 对地幔源区特征和岩石成因的约束[J]. 现代地质, 34(3): 450~465.
- 周宾, 郑有业, 许荣科, 等. 2013. 青海柴达木山岩体LA-ICP-MS锆石U-Pb定年及Hf同位素特征[J]. 地质通报, 32(7): 1 027~1 034.
- 庄玉军, 娄平阳, 李培庆, 等. 2019. 柴北缘构造带欧龙布鲁克地块西北缘辉长岩脉地球化学、年代学及Hf同位素特征[J]. 地质通报, 38(11): 1 801~1 812.