

赞皇变基性岩中锆石的 U-Pb 定年及其地质意义

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摘要: 赞皇变质杂岩区位于阜平杂岩南部, 地处华北克拉通中部造山带的中段, 和中部带北段杂岩一样, 是洞悉华北克拉通前寒武纪基底构造演化历史的一个重要窗口。研究区变基性岩可分为斜长角闪岩和角闪斜长片麻岩两种, 二者均以似层状方式产于黑云斜长片麻岩或长英质片麻岩中, 斜长角闪岩亦可呈透镜状, 二者后期与围岩一起共同经历了高角闪岩相变质作用。斜长角闪岩和角闪斜长片麻岩中普遍存在变质锆石, 锆石 SIMS U-Pb 原位定年获得的²⁰⁷Pb/²⁰⁶Pb 谐和年龄表明, 赞皇变基性岩记录了约 18.5 亿年(1 842 ± 21 Ma, 1 868 ± 29 Ma) 的一次较为广泛的变质作用事件, 结合赞皇变质杂岩的构造背景和变质演化特征, 推测该期变质作用事件与古元古代末期华北克拉通东部陆块和西部陆块间的俯冲碰撞作用有关。

关键词: 斜长角闪岩, 角闪斜长片麻岩, 锆石 U-Pb 定年, 赞皇, 华北中部带

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Zircon U-Pb dating of metabasic rocks in the Zhanhuang metamorphic complex and its geological significance

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Abstract: The Zhanhuang metamorphic complex is significant for understanding the Precambrian tectonics and evolution of the North China Craton. Metabasic rocks composed of amphibolite and amphibole-bearing plagioclase gneiss as thin layers or lenses are widely distributed within felsic gneisses and biotite-plagioclase gneisses. These metabasic rocks together with country rocks underwent high-amphibolite facies metamorphism, and abundant metamorphic zircons were formed during this thermal event. SIMS U-Pb chronological analysis indicates that metamorphic zircons record peak high-amphibolite facies metamorphism at 1 868~1 842 Ma. These Paleoproterozoic metamorphic ages are largely in accordance with metamorphic ages of ~1 850 Ma produced from high-pressure granulites in the northern segment of the Trans-North China Orogen, indicating that a significant subduction-collision event did occur in the Trans-North China Orogen between the eastern block and the western block during Paleoproterozoic (~1 850 Ma).

Key words: amphibolite; amphibole-bearing plagioclase gneiss; zircon U-Pb; Zhanhuang; Trans-North China Orogen

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关于华北克拉通早前寒武纪基底的形成和大地构造演化问题,一直以来备受国内外地质学界的关注,多个学者提出了不同的基底划分方案(白瑾等,1993,1996;沈其韩等,1995;Wang and Mo,1996;伍家善等,1998;吴昌华等,1998;张福勤等,1998;Zhao et al.,1998,2001,2005;邓晋福等,1999;李江海等,2000;翟明国等,2000;赵国春等,2002;Kusky and Li,2003;Faure et al.,2007a),且在基底的拼合机制和形成时代等方面存在较大分歧(Zhao et al.,1998,2001,2005;翟明国等,2000;Zhai et al.,2000,2005,2007;Kusky and Li,2003;Zhai and Liu,2003;Faure et al.,2007a,2007b;Kusky et al.,2007;Trap et al.,2007,2008,2009a,2009b,2011;Wang,2009;Wang et al.,2010),尤其是在前寒武纪基底的最终固化时间方面,存在~25亿年和~18.5亿年之争(Zhao et al.,1998,2001,2005;翟明国等,2000;Zhai et al.,2000,2005,2007,2010;Zhai and Liu,2003;Zhai and Santosh,2011)。位于克拉通中部、近南北向的华北克拉通中部造山带(图1 Zhao et al.,1998,2005),其构造属

性、变质演化及变质作用过程年代学对全面认识华北克拉通基底形成历史至关重要。前人对华北克拉通中部造山带北段杂岩(如建平、冀北、宣化、怀安、恒山、阜平、吕梁杂岩等)的研究比较透彻(Liu et al.,2000;Guo et al.,2002,2005a,2005b;刘树文等,2002,2009;Zhao et al.,2002b,2006,2008,2010a,2010b;Kröner et al.,2005,2006;Faure et al.,2007a,2007b;Trap et al.,2007,2008,2009a,2011;Wan et al.,2011;Yin et al.,2011),认为其普遍经历了高角闪岩相-麻粒岩相变质作用,拥有顺时针、退变质阶段为近等温降压(ITD)型变质作用P-T轨迹,记录了18.5亿年左右的变质事件,卷入了华北克拉通中部造山带的碰撞造山运动。相对而言,在中部带中南段的变质杂岩区(如赞皇、中条、登封、太华变质杂岩等),虽然一些学者也在该地区进行过许多很好的工作(涂绍雄,1984,1998;齐进英,1991,1992;薛良伟等,1995,2005;白瑾等,1997;周汉文等,1997,1998;倪志耀等,2003;王岳军等,2003;Wang et al.,2003b,2004;Tian et al.,2006;Wan et al.,2006,2011;赵凤清,2006;刘树文等,

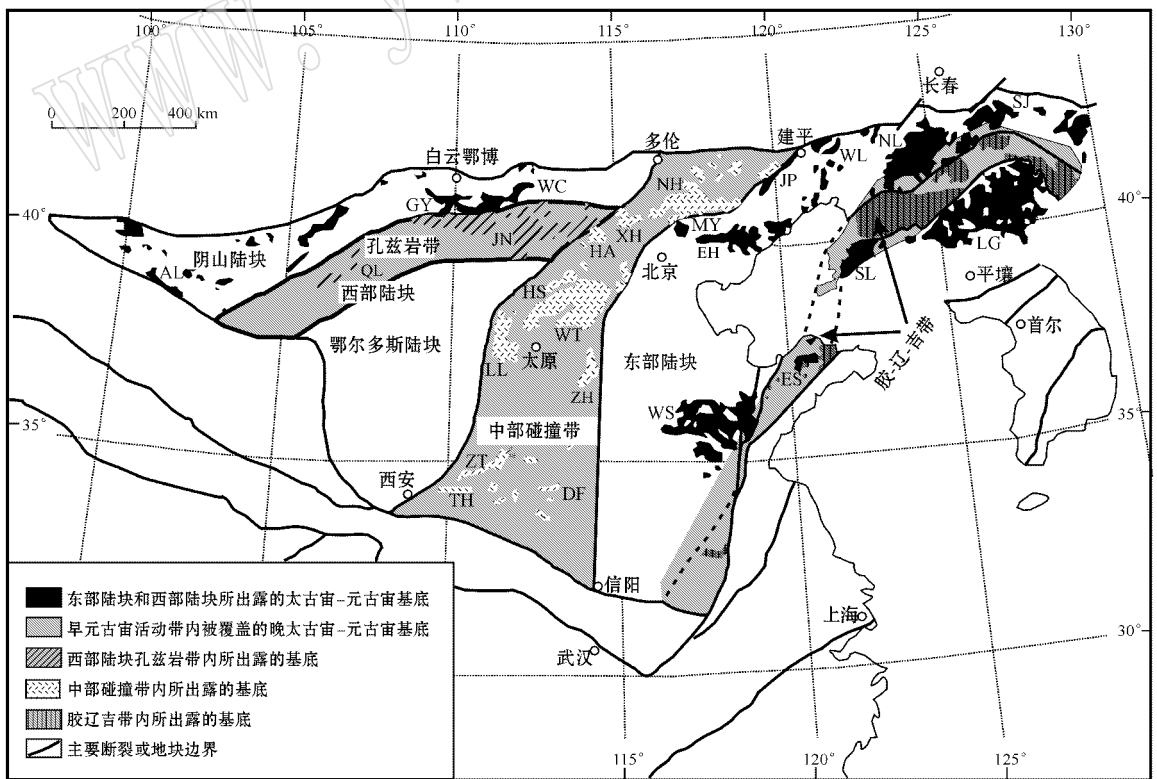


图1 华北克拉通前寒武纪基底构造划分图(据 Zhao et al., 1998, 2001, 2005)

Fig.1 Geological sketch map of the North China Craton(after Zhao et al., 1998, 2001, 2005)

2007, 2009; 吴春明等, 2008; 肖玲玲等, 2008, 2010, 2011; 杨长秀, 2008; He *et al.*, 2009; Trap *et al.*, 2009a; 万渝生等, 2009; Xu *et al.*, 2009; Diwu *et al.*, 2010; Xiao *et al.*, 2011; 杨崇辉等, 2011), 但这些杂岩区在变质演化和变质作用年代学方面的研究还比较薄弱, 尤其是位于华北克拉通中部带中段的赞皇变质杂岩, 其在变质作用年代学方面的报道资料相对较少, 亟待深入研究。赞皇变质杂岩是否与北段其他杂岩一样, 经历了 18.5 亿年左右的变质事件, 卷入了华北中部带的碰撞造山作用, 这些问题制约了我们对华北中部带的属性和构造演化完整过程的认识。

赞皇变质杂岩区出露的变基性岩普遍经历了高角闪岩相峰期变质作用(肖玲玲等, 2011; Xiao *et al.*, 2011), 形成于类似现代岛弧的构造环境(肖玲玲等, 2010), 岩石所记录的年代学信息能够较好地反映赞皇变质杂岩形成以来所经历的构造演化历程, 有助于更全面地认识华北克拉通中部造山带前寒武纪基底构造演化过程。前人已在该地区的变质演化、构造及年代学等方面做过一些研究工作(马文璞等, 1989; 雷世和, 1991; 雷世和等, 1994; 牛树银, 1994; 牛树银等, 1994a, 1994b; 王岳军等, 2003; Wang *et al.*, 2003b, 2004; 肖玲玲等, 2008, 2010, 2011; Trap *et al.*, 2009a; Xiao *et al.*, 2011; 杨崇辉等, 2011), 为本文的工作提供了很好的研究基础。本文在对赞皇地区出露的两类变基性岩(斜长角闪岩和角闪斜长片麻岩)进行详细的野外地质调查、室内岩相学研究的基础上, 分别对其典型样品进行锆石的离子探针(SIMS)测年, 以获得相应变质过程的年代学信息。

1 地质背景

赞皇变质杂岩区主体出露于河北省境内, 总体呈一中部宽、两端收拢的“纺锤”形, 南北向展布, 东西宽 40~60 km, 南北长约 140 km, 长轴北端近南北, 南端呈 NNE-SSW 走向, 总面积约 4 000 km²(图 2, 河北省地质局^①, 牛树银, 1994)。赞皇变质杂岩区地处太行山山脉中段东麓, 构造位置上处于华北克拉通中部造山带的中段(图 1, Zhao *et al.*, 1998, 2001, 2005), 向北紧邻阜平杂岩区。该地区出露有

较多的前寒武纪岩石, 其中以太古界岩石出露最多、分布面积最大^①(图 2), 是研究华北克拉通东、西部陆块碰撞造山作用的重要地区之一。从岩性分布上看, 区内出露的主要变质岩石类型有 TTG 片麻岩、长英质片麻岩、(含榴蓝晶)黑云斜长片麻岩、(含榴)斜长角闪岩、角闪斜长片麻岩、磁铁石英岩、大理岩、变质砂岩等, 其中出露最多的为 TTG 片麻岩, 其他各类变质岩石出露面积仅在 10% 以内, 零星分布于赞皇变质杂岩区的中部和南部地区(图 2, 杨崇辉等, 2011)。

赞皇变质杂岩区内部变形作用比较复杂, 广泛发育倒转褶皱、横弯褶皱、平卧褶皱等(牛树银等, 1994a), 区内可识别出 3 条 NNE 向韧性剪切带, 自西向东分别为: 营房台-招七-障石岩-苍岩山剪切带、邢台坡底-临城郝庄-官都剪切带、赞皇榆底村-临城岗西-元氏黑水河剪切带(牛树银等, 1994a, 1994b; 王岳军等, 2003; Wang *et al.*, 2003b)。雷世和(1991)、牛树银等(1994a, 1994b)认为赞皇变质杂岩区内 3 条拆离韧性剪切带形成于中生代, 受燕山运动影响, 该杂岩于中生代以伸展体制为主的构造环境当中发生快速隆起, 形成一个典型的变质核杂岩区——赞皇变质核杂岩(马文璞等, 1989; 雷世和等, 1994; 牛树银, 1994; 牛树银等, 1994a, 1994b), 早中元古界及其以上地层构成该变质核杂岩的盖层拆离滑脱系(牛树银, 1994), 但未见相关年龄数据支持该观点。河北省地质矿产局(1989)和王岳军等(王岳军等, 2003; Wang *et al.*, 2003b, 2004)对该地区进行了详细的区域地质调查及构造历史研究, 并利用黑云母⁴⁰Ar/³⁹Ar 年代学方法, 对赞皇变质杂岩区主要热事件年龄和韧性剪切带的形成时代进行了厘定, 认为赞皇变质杂岩区的形成是早元古代晚期-中元古代多次构造热事件作用的结果, 并不是一个简单的中生代变质核杂岩, 而是早元古代末期形成的变质穹窿(王岳军等, 2003)。

2 野外产状及岩相学特征

赞皇斜长角闪岩出露于赞皇变质杂岩的中南部地区, 以“似层状”(图 3a, 3b)或“透镜状”(图 3c)产出, 于黑云斜长片麻岩或长英质片麻岩中, “似层状”斜长角闪岩局部被拉(断)成“石香肠”构造(图 3a)或发

① 河北省地质局, 1968. 1:200 000 高邑幅、邢台幅地质图。

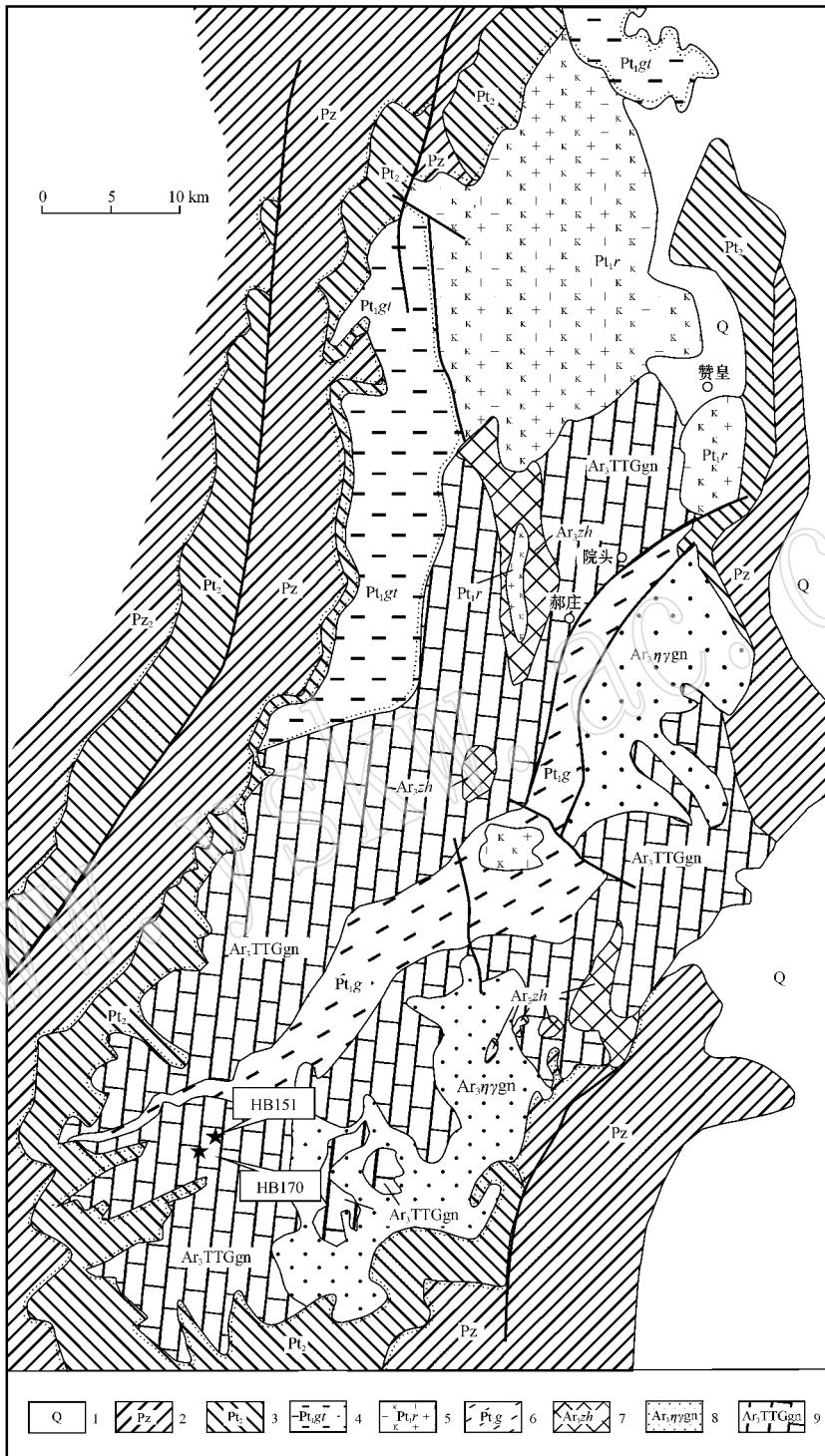


图 2 赞皇变质杂岩地质简图及采样点位置(据杨崇辉等, 2011)

Fig. 2 Geological sketch map of the Zanzhuang Metamorphic Complex, showing representative sampling locations (after Yang Chonghui *et al.*, 2011)

1—第四系; 2—古生代地层; 3—中元古代地层; 4—古元古代甘陶河群; 5—许亭花岗岩; 6—古元古代官都群; 7—新太古代赞皇群; 8—新太古代二长花岗岩片麻岩; 9—新太古代 TTG 片麻岩

1—Quaternary; 2—Paleozoic; 3—Mesoproterozoic Changcheng Group; 4—Paleoproterozoic Gantaohu Group; 5—Xuting Granite; 6—Paleoproterozoic Guandu Group; 7—Neoproterozoic Zanzhuang Group; 8—Neoproterozoic Monzogranite gneiss; 9—Neoproterozoic TTG gneiss

生明显褶皱变形(图 3b)。斜长角闪岩中通常含有石榴石,且石榴石变斑晶周围多发育“白眼圈”反应结构,指示快速降压过程(肖玲玲等 2011; Xiao *et al.*, 2011)。基质矿物主要包括角闪石、单斜辉石、斜长石、石英及少量黑云母、榍石、磁铁矿、金红石、钛铁矿、黄铁矿、锆石等。赞皇地区斜长角闪岩不同阶段矿物组合及其变形关系如图 4 所示。

本文斜长角闪岩样品 HB151 采自邢台市邢台县城计头乡花木村,其岩相学特征详细描述如下,矿物缩写据沈其韩(2009)。含榴斜长角闪岩 HB151: 灰黑色,斑状变晶结构,单斜辉石变斑晶粒径可达 3 mm。石榴石颗粒较小,直径约 0.1~0.2 mm,呈港湾状或残块状(图 5a),片麻状构造,可识别出两个阶段的变质矿物组合:①峰期变质矿物组合,由单斜辉石变斑晶和基质石榴石、角闪石、斜长石、石英及少量黑云母、磁铁矿、磷灰石、锆石等组成,轴向矿物的定向排列构成了宏观片麻理;②退变质阶段矿物组

合,表现为围绕单斜辉石变斑晶边部发育的绿色“镶边状”角闪石(图 5b)或在变斑晶内部呈定向分布的细条状角闪石(图 5b),二者均由峰期矿物单斜辉石分解生成。

赞皇角闪斜长片麻岩主体分布于赞皇变质杂岩的中北部地区,多以厚层状产出于长英质片麻岩中(图 3d)。采自邢台市邢台县路罗镇大戈寥村的角闪斜长片麻岩样品 HB170 呈青灰色,中细粒变晶结构,片麻状构造,且片麻理非常明显,组成矿物主要包括角闪石、斜长石、石英、磁铁矿及少量榍石、金红石、锆石等(图 5d),为峰期变质矿物组合。

前期地球化学和变质演化研究表明,赞皇变基性岩形成于类似现代岛弧构造环境(肖玲玲等, 2010),普遍经历了早期前进变质、峰期高角闪岩相变质、峰后的角闪岩相退变质作用过程(肖玲玲等, 2008, 2011; Xiao *et al.*, 2011)。结合野外产状推测赞皇斜长角闪岩的原岩为基性侵入岩墙或岩脉,角

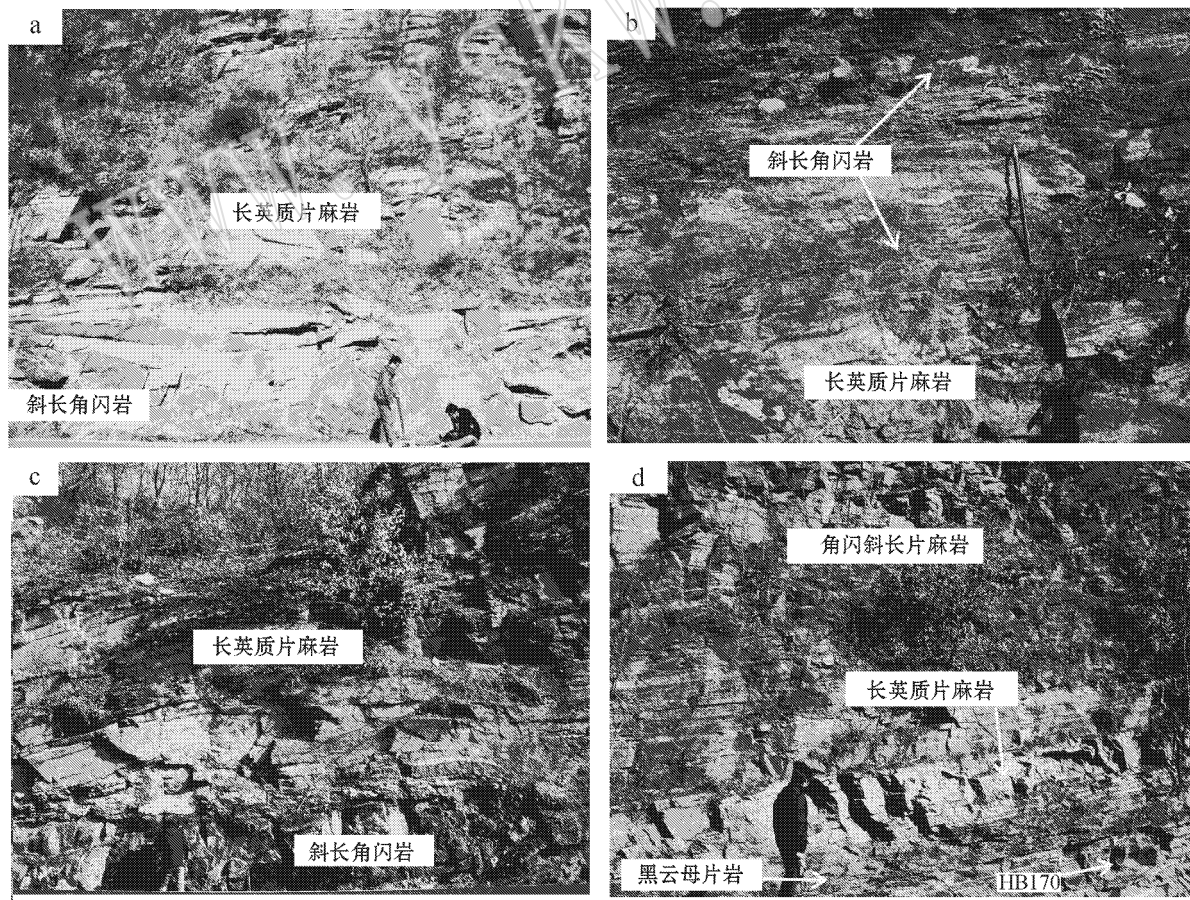


图 3 “似层状”(a、b)“透镜状”(c)产出的斜长角闪岩以及“厚层状”(d)产出的角闪斜长片麻岩

Fig. 3 Stratoid(a、b) and lenticular(c) amphibolites, and stratoid amphibole-bearing plagioclase gneiss(d)

变形期次	D ₁	D ₂	D ₃ (?)
变质期次	进变质作用阶段 M ₁	近峰期变质作用阶段 M ₂	退变质作用阶段 M ₃
矿物			
石榴石	■	■	
角闪石	■	■	
斜长石	■	■	
石英	■	■	
单斜辉石		■	
黑云母	■	■	
钾长石	■	■	
楣石	■	■	
磷灰石	■	■	
磁铁矿	■	■	
钛铁矿	■	■	
黄铁矿		■	
金红石		■	
锆石	■	■	

图 4 赞皇变质性岩不同阶段矿物组合及其变形关系图

Fig. 4 Mineral crystallization-deformation diagram of Zanhuang metabasic rocks

闪斜长片麻岩的原岩为基性火山岩层。

3 锆石 U-Pb 定年

锆石的 U-Pb 同位素体系具有非常高的封闭温度 (Lee *et al.*, 1997; Cherniak and Watson, 2000), 目前锆石 U-Pb 定年已成为放射同位素年代学研究中最为常用和最有效的方法之一。本文对赞皇变质杂岩区中斜长角闪岩样品 HB151 和角闪斜长片麻岩样品 HB170 进行了锆石 SIMS U-Pb 年代学测试。

3.1 测试方法

阴极发光 (CL) 图像拍摄在中国科学院地质与地球物理研究所离子探针中心进行, 用于细致观察锆石内部结构, 如有无环带、有无裂缝、有无核边结构等, 结合透、反射光图像, 选择年代学实验的待测位置、分析锆石成因。锆石的 SIMS U-Pb 定年在该离子探针中心的 CAMECA 1280 仪器上测试完成, 测年方法和原理参见 Li 等 (2009)。样品数据点 (包括标准锆石) 测定为 5 组扫描, 一次离子流强度为 6 nA, 束斑直径为 30 μm 。U 含量和同位素比值分别

采用标准锆石 91500 (81×10^{-6} , Wiedenbeck *et al.*, 1995) 和 Plésovice (337 Ma, Sláma *et al.*, 2008) 校正获得, 测定锆石标样与锆石样品的比例约为 1:3。分别运用 SQUID 和 ISOPLOT 程序 (Ludwig, 2003) 进行数据处理和谐和曲线绘制。同位素比值和单个年龄实测误差为 1σ , 加权平均年龄误差为 95% 置信度。

3.2 测试结果

本文赞皇斜长角闪岩和角闪斜长片麻岩样品中的锆石颗粒普遍较小, 粒径多为 10~90 μm , 个别颗粒可达 120 μm , 形态不规则, 多呈浑圆状或短柱状, 浅黄色-浅色, 透明, 边缘模糊。阴极发光 (CL) 图像显示 (图 6a、6e), 锆石不发育核-边结构, 无韵律环带, 内部结构不均匀, 呈斑杂状或云雾状, 显示变质锆石特征。锆石 U-Pb 同位素测试结果见表 1。

含榴斜长角闪岩 HB151: 对该样品中 19 颗锆石的 19 个点分析表明 (表 1), 锆石 Th 含量介于 $0.01 \times 10^{-6} \sim 64.2 \times 10^{-6}$, U 含量介于 $0.9 \times 10^{-6} \sim 140.2 \times 10^{-6}$, Th/U 比值为 0.001~0.458, 其中 16 个测试点的 Th/U 比值 < 0.1 , Th、U 含量无明显相

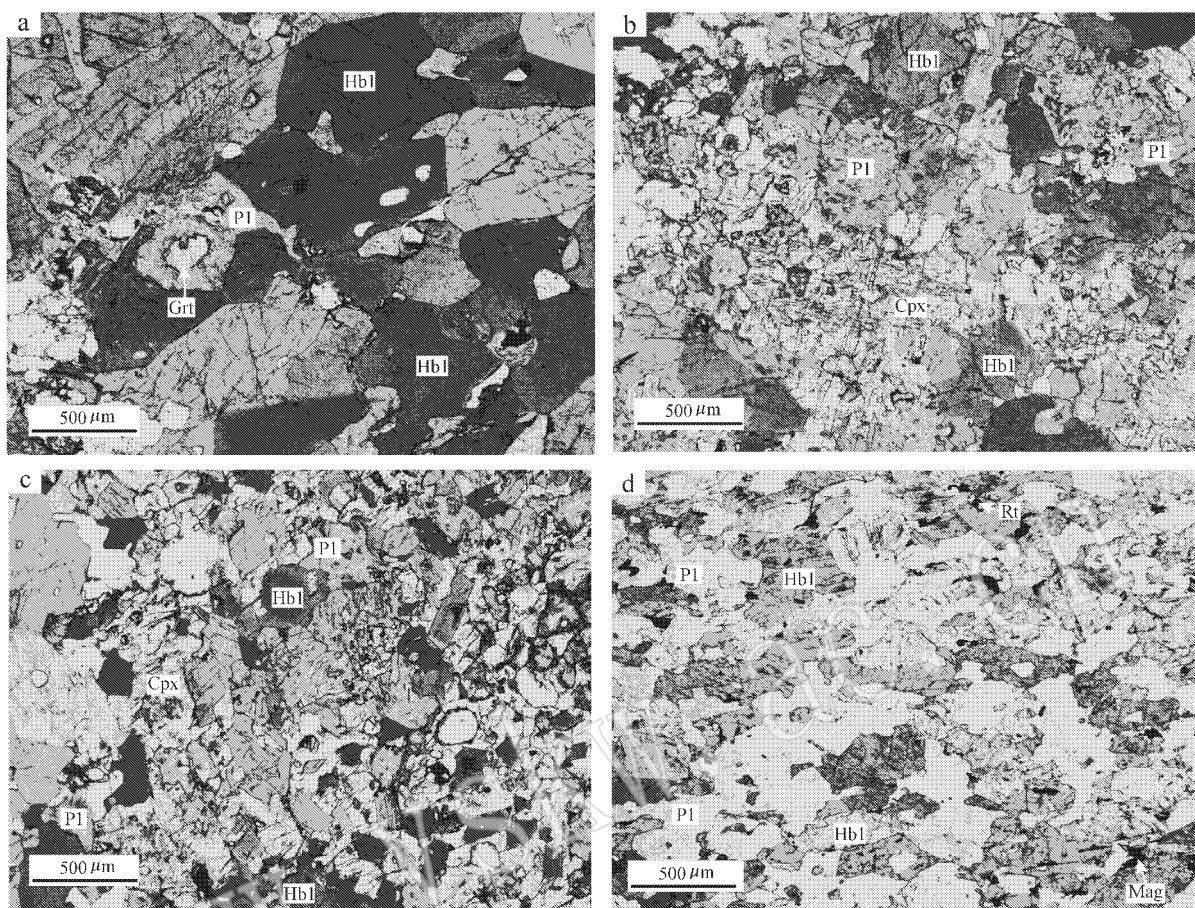


图 5 赞皇斜长角闪岩样品 HB151(a~c)和角闪斜长片麻岩样品 HB170(d)的单偏光镜下的显微岩相照片

Fig. 5 Polarizing optical photomicrographs of Sample HB151 (a~c) and Sample HB170 (d)

关性(图 6d),显示变质成因锆石特征。样品 HB151 中 15 个有效数据点的谐和年龄为 $1\ 842 \pm 21\ \text{Ma}$ (MSWD = 1.4), 加权平均年龄为 $1\ 841 \pm 16\ \text{Ma}$ (MSWD = 1.4)(图 6b、6c)。

角闪斜长片麻岩 HB170: 对其中 14 颗锆石进行了 14 个点分析(表 1), 结果表明, 锆石 Th 含量介于 $0.02 \times 10^{-6} \sim 23.2 \times 10^{-6}$, U 含量变化较大 ($2.5 \times 10^{-6} \sim 1\ 408.9 \times 10^{-6}$), Th/U 比值为 $0.006 \sim 0.016$, 所有分析点 Th/U 比值均 < 0.1 , 与阴极发光图像一样, 均显示变质成因锆石特征。除 HB170-11 外, 其余各分析点的 Th、U 含量趋于正相关(图 6h)。样品 HB170 中 11 个有效数据点的谐和年龄为 $1\ 868 \pm 29\ \text{Ma}$ (MSWD = 1.16), 10 个有效数据点的加权平均年龄为 $1\ 850 \pm 23\ \text{Ma}$ (MSWD = 0.59)(图 6f、6g)。

3.3 年代学结果探讨

赞皇斜长角闪岩和角闪斜长片麻岩中普遍发育变质锆石, 本文所获得的两个样品的变质锆石 $^{207}\text{Pb}/$

^{206}Pb 年龄分别为 $1\ 842 \pm 21\ \text{Ma}$ 和 $1\ 868 \pm 29\ \text{Ma}$, 与赞皇变泥质岩所记录的变质年龄 ($1\ 821 \pm 17\ \text{Ma}$) 近于一致(肖玲玲等, 2011)。前人在该地区也做过一些年代学工作: Wang 等(2003b)在该地区辨认出 4 期变形作用, 应用单矿物 $^{40}\text{Ar}/^{39}\text{Ar}$ 定年方法得到各期变形作用年龄分别为: $1\ 870\ \text{Ma}$ (D_1), $1\ 870 \sim 1\ 826\ \text{Ma}$ (D_2), $1\ 826 \sim 1\ 793\ \text{Ma}$ (D_3) 和 $1\ 689 \sim 1\ 633\ \text{Ma}$ (D_4), 并认为赞皇变质杂岩区的形成, 是早元古代晚期—中元古代多次构造事件共同作用的结果, 影响赞皇变质杂岩区的主要构造事件发生在 $1\ 870 \sim 1\ 793\ \text{Ma}$ 。Trap 等(2009a)将赞皇变质杂岩区划分为西部区块、中部区块和东部区块, 在中部区块西缘存在一条构造缝合带, 推测赞皇变质杂岩区的形成是 $1\ 880 \sim 1\ 850\ \text{Ma}$ 期间华北克拉通东部陆块与阜平陆块沿构造缝合带碰撞拼合作用的结果。同时, Trap 等(2009a)对赞皇地区的黑云斜长片麻岩、斜长角闪岩和正片麻岩分别进行了独居石 EMPA

表 1 赞皇变基性岩样品 HB151 和 HB170 的锆石 SIMS U-Pb 分析数据

Table 1 SIMS U-Pb zircon data of metabasic samples HB151 and HB170 in the Zhanhuang metamorphic complex

分析点号	$\omega_{\text{Pb}}/10^{-6}$		Th/U	$^{207}\text{Pb}/^{235}\text{U}$	1 σ	$^{206}\text{Pb}/^{238}\text{U}$	1 σ	$^{208}\text{Pb}/^{232}\text{Th}$	1 σ	年龄/Ma					
										$^{207}\text{Pb}/^{206}\text{Pb}$		$^{207}\text{Pb}/^{235}\text{U}$		$^{206}\text{Pb}/^{238}\text{U}$	
	Th	U								1 σ	1 σ	1 σ	1 σ		
HB151-1	0.25	13.83	0.018	4.527 40	0.038 87	0.294 03	0.022 76	0.105 24	0.251 83	1 827	56	1 736	33	1 662	33
HB151-2	1.99	12.53	0.159	4.177 96	0.027 78	0.271 08	0.016 10	0.049 58	0.162 37	1 829	40	1 670	23	1 546	22
HB151-3	0.29	1.52	0.191	5.445 21	0.090 10	0.298 05	0.015 76	0.116 98	0.372 57	2 131	148	1 892	80	1 682	23
HB151-4	0.88	24.83	0.035	4.579 00	0.020 59	0.298 74	0.015 47	0.061 45	0.248 48	1 819	24	1 745	17	1 685	23
HB151-5	0.07	3.67	0.019	5.210 42	0.036 90	0.310 04	0.019 01	0.250 60	0.147 22	1 984	55	1 854	32	1 741	29
HB151-6	0.19	0.93	0.201	5.095 56	0.087 27	0.308 32	0.018 77	0.109 39	0.147 61	1 954	145	1 835	77	1 732	29
HB151-8	0.25	27.02	0.009	4.815 71	0.019 15	0.311 78	0.015 31	0.095 20	0.128 14	1 833	21	1 788	16	1 749	23
HB151-9	0.10	4.41	0.023	5.071 48	0.037 44	0.322 55	0.017 44	0.029 38	2.715 99	1 865	59	1 831	32	1 802	27
HB151-10	0.19	25.44	0.007	5.067 90	0.020 16	0.319 91	0.015 16	0.099 29	0.141 87	1 878	24	1 831	17	1 789	24
HB151-11	0.12	15.76	0.007	4.863 80	0.021 23	0.322 33	0.015 27	0.147 13	0.156 08	1 790	27	1 796	18	1 801	24
HB151-12	64.21	140.24	0.458	11.984 96	0.015 53	0.497 10	0.015 17	0.142 94	0.048 33	2 605	6	2 603	15	2 601	33
HB151-13	0.20	23.31	0.009	4.940 22	0.019 24	0.320 36	0.015 01	0.099 07	0.134 92	1 830	22	1 809	16	1 791	24
HB151-14	0.23	10.16	0.022	5.285 64	0.024 10	0.342 48	0.015 79	0.114 89	0.126 28	1 831	33	1 867	21	1 899	26
HB151-15	0.09	27.28	0.003	5.116 14	0.019 73	0.331 58	0.016 32	0.147 09	0.172 75	1 831	20	1 839	17	1 846	26
HB151-16	0.01	7.64	0.001	5.206 83	0.030 17	0.329 40	0.019 80	0.880 93	0.313 68	1 874	40	1 854	26	1 835	32
HB151-17	0.03	12.70	0.002	5.293 41	0.027 96	0.328 10	0.015 03	0.286 41	0.273 25	1 911	42	1 868	24	1 829	24
HB151-18	0.24	2.77	0.085	4.876 53	0.044 02	0.304 73	0.015 74	0.106 63	0.145 27	1 896	72	1 798	38	1 715	24
HB151-19	0.13	3.22	0.040	4.702 41	0.046 28	0.295 43	0.026 46	0.122 57	0.177 90	1 887	67	1 768	40	1 669	39
HB170-1	0.50	45.82	0.011	4.749 077	0.018 498	0.306 811	0.015 060	0.098 720	0.099 812	1 836	19	1 776	16	1 725	23
HB170-2	0.12	17.45	0.007	4.851 131	0.021 129	0.308 513	0.015 258	0.159 224	0.205 987	1 865	26	1 794	18	1 733	23
HB170-3	0.09	6.99	0.014	4.807 442	0.038 373	0.307 852	0.015 258	0.132 113	0.170 625	1 852	62	1 786	33	1 730	23
HB170-4	0.02	2.50	0.008	4.559 829	0.126 348	0.299 995	0.015 076			1 803	212	1 742	111	1 691	22
HB170-5	0.02	3.29	0.007	4.617 285	0.046 024	0.303 041	0.016 165			1 808	76	1 752	39	1 706	24
HB170-6	0.17	27.53	0.006	4.944 548	0.023 007	0.311 539	0.015 144	0.272 254	0.573 273	1 882	31	1 810	20	1 748	23
HB170-7	0.07	10.55	0.007	4.826 006	0.025 455	0.305 760	0.016 281	0.324 653	0.355 907	1 872	35	1 789	22	1 720	25
HB170-8	0.06	6.09	0.009	4.668 974	0.043 499	0.302 419	0.015 625	0.284 121	0.160 301	1 832	72	1 762	37	1 703	23
HB170-9	0.12	16.48	0.007	4.679 743	0.032 396	0.312 381	0.015 202			1 777	51	1 764	27	1 752	23
HB170-10	0.04	4.56	0.008	4.961 565	0.070 699	0.300 060	0.015 026	1.352 472	0.563 928	1 955	118	1 813	62	1 692	22
HB170-11	23.21	1 408.86	0.016	1.765 273	0.015 474	0.139 212	0.015 004	0.041 906	0.150 329	1 467	7	1 033	10	840	12
HB170-12	0.25	24.85	0.010	4.389 075	0.042 684	0.309 002	0.015 324			1 679	72	1 710	36	1 736	23
HB170-14	0.05	4.67	0.010	4.794 932	0.037 975	0.304 795	0.020 173	0.235 052	0.205 562	1 866	57	1 784	32	1 715	30

Th-U-Pb 和角闪石 $^{40}\text{Ar}/^{39}\text{Ar}$ 定年,其中,黑云斜长片麻岩基质中的独居石形成于 $1\,824 \pm 6$ Ma,角闪岩和正片麻岩得到的角闪石 $^{40}\text{Ar}/^{39}\text{Ar}$ 年龄约为 1 800 Ma。综合以上,可以看出 $\sim 1\,850$ Ma 的变质事件在赞皇变质杂岩区记录非常普遍,可能代表了该地区所经历的一次大规模的构造事件。

4 大地构造意义

目前,关于华北克拉通早前寒武纪基底的划分方案很多(白瑾等,1993,1996;沈其韩等,1995;Wang and Mo,1996;伍家善等,1998;吴昌华等,1998;张福勤等,1998;Zhao *et al.*,1998,2001,2005;邓晋福等,1999;李江海等,2000;翟明国等,

2000;赵国春等,2002;Kusky and Li,2003;Faure *et al.*,2007a),且在基底的具体碰撞方式和时间等方面亦存在明显分歧:①翟明国等(翟明国等,2000;Zhai *et al.*,2000,2005,2007;Zhai and Liu,2003)通过识别重要的古老构造边界,将华北克拉通前寒武纪基底划分为 6 个微陆块(胶辽、迁怀、阜平、集宁、许昌、阿拉善),各微陆块普遍经历了新太古代末期的变质-构造-热事件,于 2.5 Ga 左右彼此拼合形成了华北克拉通前寒武纪结晶基底,具体拼贴模式尚待进一步研究。随后华北克拉通处于相对稳定期, ~ 1.8 Ga 时基底已发生大规模裂解;②Zhao 等(1998,2001,2005)根据变质作用 P - T 轨迹形状的不同,将华北克拉通早前寒武纪变质基底划分为 3 个部分(图 1):东部陆块(the Eastern Block)、西部陆

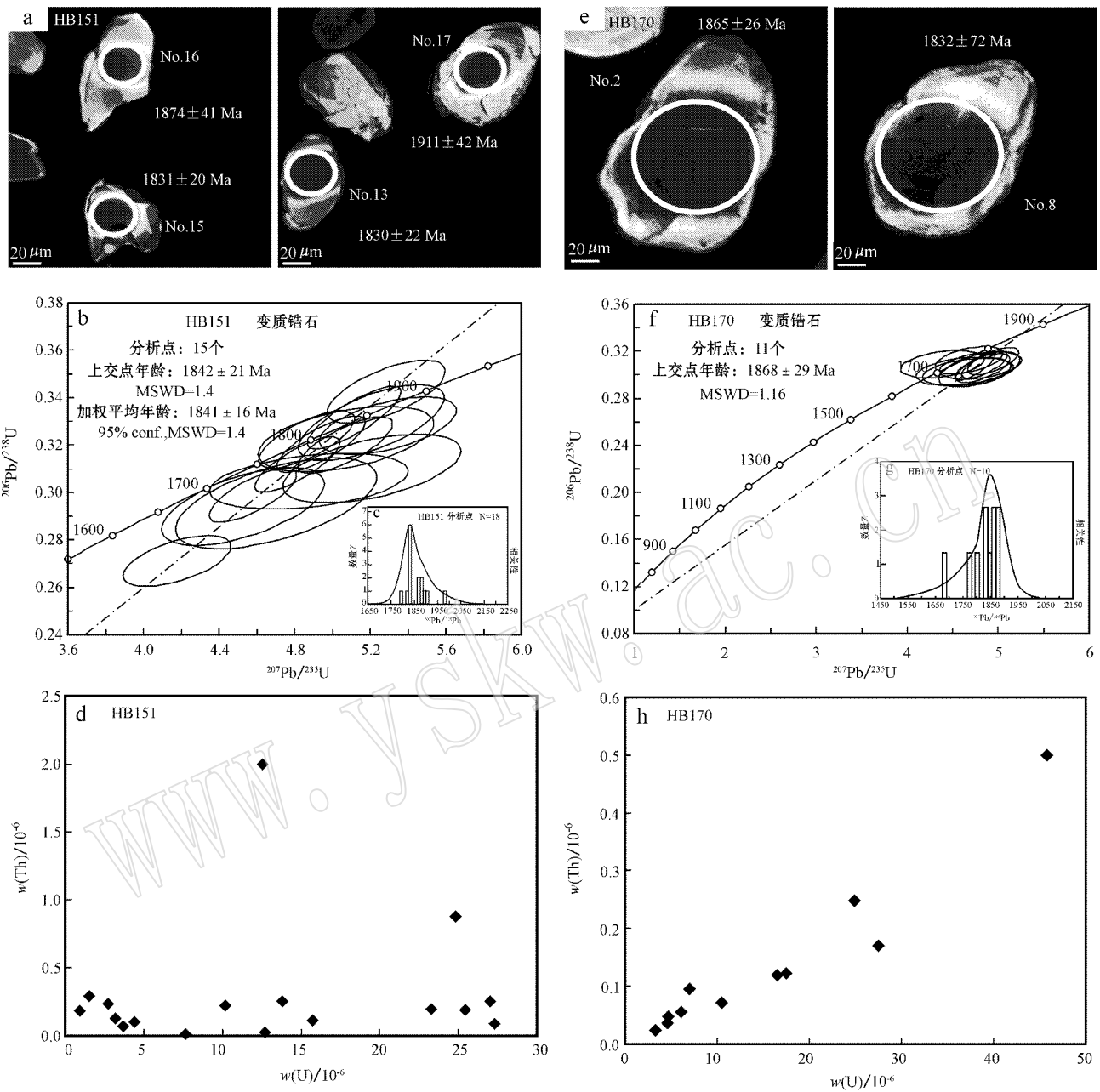


图 6 赞皇变基性岩的锆石阴极发光图像 (a, e)、SIMS 锆石 U-Pb 谐和曲线图 (b, f)、锆石年龄频度分布图 (c, g) 以及锆石 U-Th 含量相关关系图 (d, h) (用于标准化的 C1 球粒陨石组成据 Sun and McDonough, 1989)

Fig. 6 Representative cathodoluminescence images (a, e), and U-Pb concordia diagram, showing SIMS analytical data (b, f), distribution map of age frequency (c, g) and U-Th diagram (d, h) for zircon of amphibolite from the Zhanhuang metamorphic complex (chondrite 1-normalized data from Sun and McDonough, 1989)

块 (the Western Block) 和华北中部造山带 (the Trans-North China Orogen), 认为统一的华北克拉通基底是 ~1.85 Ga 时由东部和西部两个陆块沿近南北向的“华北中部造山带” (the Trans-North China Orogen) 碰撞拼合而成; ③Kusky 和 Li (2003) 亦有华北

克拉通前寒武纪基底三分的构造模式, 但将东、西部陆块之间的碰撞带定义为“中央造山带” (the Central Orogenic Belt), 其构造边界与赵国春等划分的“华北中部造山带” (the Trans-North China Orogen) 略有不同 (Kusky and Li, 2003; Kusky *et al.*, 2007)。同时

认为,东部地块与西部地块是在~2.5 Ga 以弧-陆形式发生汇聚,~1.8 Ga 以后华北克拉通以伸展构造体制为主,期间形成一系列拗拉槽和裂谷,并伴有大量基性岩浆的侵入;④Faure 等(2007a)将阜平微陆块单独划分出来,认为华北克拉通太古宙基底总体可分为鄂尔多斯陆块、阜平陆块、东部陆块3个部分,三陆块之间分别被一个古洋盆“吕梁洋”和“太行洋”所分隔。整个华北克拉通前寒武纪基底的聚合,是由两个阶段沿中部带发生的東西向碰撞事件完成的(Faure *et al.*, 2007a, 2007b; Trap *et al.*, 2007, 2008, 2009a, 2009b, 2011): ~2.1 Ga 东部陆块发生西向俯冲,首先与阜平陆块发生碰撞,导致“太行洋”的关闭和一系列岛弧岩浆作用的发生,而后继续向西俯冲,在1.9~1.8 Ga 时,“吕梁洋”闭合,阜平陆块与西部陆块碰撞拼合,最终形成统一的华北克拉通基底;⑤Wang 等(Wang, 2009; Wang *et al.*, 2010)认为华北克拉通东、西部陆块之间存在一个古华北洋(>~2 565 Ma, the Proto-North China Ocean),~2 565—2 540 Ma 北西向的洋内俯冲形成了以大量花岗岩岩石为代表的五台原始岛弧(proto-arc),随后~2 525—2 475 Ma 期间多个俯冲带的多阶段俯冲-增生作用,导致了五台弧后盆地(the Wutai back-arc basin)的关闭和大量不同类型的岛弧岩浆的形成。虽然对~2.5—1.8 Ga 之间发生的事情尚不确定,但Wang(2009)认为华北克拉通东、西部陆块于~1.9—1.8 Ga 发生了大规模的碰撞造山作用,最终导致了华北中部造山带和华北克拉通前寒武纪基底的形成。

综上所述,目前对华北克拉通前寒武纪基底构造格架的研究,虽然在某些划分细节和拼合时间上有所差别,但已经形成一个基本共识,即华北克拉通前寒武纪基底是由先存的数个微陆块后期拼贴而成的。并且越来越多的年代学数据(SHRIMP 锆石 U-Pb、矿物⁴⁰Ar/³⁹Ar 和 Sm-Nd、EMPA 独居石 Th-U-Pb 年龄等)表明,华北克拉通1 900~1 800 Ma 期间的变质事件记录明显,尤其是位于华北中部造山带的各变质杂岩区(Liu *et al.*, 2000; 刘树文等, 2002, 2009; 赵国春等, 2002; Zhao *et al.*, 2002a, 2006, 2008, 2010a; Guo *et al.*, 2005a, 2005b; Kröner *et al.*, 2005, 2006; Faure *et al.*, 2007a, 2007b; Trap *et al.*, 2007, 2008, 2009a, 2011; 肖玲玲等, 2011),而这期大规模变质事件与华北克拉通东部陆块与西部陆块之间的碰撞造山作用密切相关。

与中部带其他杂岩区类似,赞皇变质杂岩区中的变基性岩和变泥质岩(肖玲玲等, 2011)同样经历了~1 850 Ma 的一期广泛的变质作用,具有顺时针近等温降压型的 $P-T-t$ 轨迹(肖玲玲等, 2011),笔者推断,赞皇变质杂岩区很可能也卷入了华北克拉通前寒武纪基底的碰撞造山运动,该区变基性岩所记录的变质时代为1 868~1 842 Ma。

5 结论

(1) 赞皇角闪斜长片麻岩中保存了两个阶段的变质作用矿物组合,包括峰期矿物组合 $Grt + Hbl + Cpx + Pl + Qtz$ 和退变质矿物组合 $Hbl + Pl + Qtz$;

(2) 赞皇斜长角闪岩和角闪斜长片麻岩中均发育变质成因锆石,同变泥质岩一样,记录了该地区所经历的一期广泛的高角闪岩相变质热事件,变质时代为1 868~1 842 Ma;

(3) 结合野外地质特征及前期地球化学和变质演化研究,推测赞皇变质杂岩曾卷入了华北克拉通早前寒武纪基底东、西两大陆块间的一次大规模的碰撞造山运动。

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