

## 桂西二叠系铝土岩系中的灾变事件记录研究

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**摘要:** 灾变事件赋存着重要的地质信息, 其遗迹具有经济和科学价值, 发现和研究灾变事件是解决地质问题的方法之一, 也是提高地质科学认识水平的途径。本文通过桂西地区铝土矿床野外调查、显微特征研究和含铝岩系重金属元素分析, 列举桂西地区二叠系含铝岩系主要灾变事件, 阐释其形成诱因。含铝岩系灾变事件主要为火山灰季、重金属异常和地震幕。火山灰降落时长跨度可达 10 Ma, 火山灰岩性从基性到酸性均有出现, 属岛弧钙碱性岩浆系列。相对于华北、黔北等铝土矿集区, 桂西含铝岩系中 Cr 强烈富集, 剖面平均值为  $400 \times 10^{-6}$  ~  $900 \times 10^{-6}$ , 在部分剖面中强烈富集, 平均值  $300 \times 10^{-6}$  ~  $600 \times 10^{-6}$ , Cd 在大多数剖面中富集, Co 在少数剖面中富集, 在酸性火山灰的岩层中 Sb、As 强烈富集, Cr、Ni、Cd、Sb、As、Co 等 6 种元素构成重金属异常。地层遗迹可见软变形、震浊积岩、小地堑等。火山灰降落和地震幕的诱因应为印支板块与华南板块的俯冲形成岛弧岩浆带火山喷发, 同时弧后盆地拉张引发了地震。火山灰不仅为含铝岩系提供了铝、铁等物质, 并且提升了重金属, 导致重金属异常。

**关键词:** 灾变事件; 含铝岩系; 火山灰; 重金属; 桂西

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## The records of catastrophic events in the Permian bauxite rock series in western Guangxi

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**Abstract:** Catastrophic events contain important geological information, and their relics have economic and scientific value. Discovering and studying catastrophic events is one of the methods to solve geological problems and also a way to improve the understanding of geological science. This article lists the main catastrophic events of the Permian bauxite rock series in western Guangxi basing on field investigations, microscopic feature studies, and analysis of heavy metal elements in the bauxite-bearing rock series, and explains their formation. The main catastrophic events in bauxite-bearing rock formations are volcanic ash season, heavy metal anomalies and earthquake episode. The duration span of volcanic ash deposition can extend to 10 Ma, and the lithology of volcanic ash varies from basic to acidic, showing the arc-related calc-alkaline magma series. Compared to bauxite ore concentration areas in the

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North China and North Guizhou, the bauxite-bearing rock series in studied sections of western Guangxi is strongly enriched in Cr, with an average value of  $400 \times 10^{-6}$  to  $900 \times 10^{-6}$ , whereas Ni is strongly enriched in some sections, with an average value of  $300 \times 10^{-6}$  ~  $600 \times 10^{-6}$ , Cd is enriched in most sections, but Co is enriched in few sections. In addition, Sb and As are greatly enriched in the rocks with felsic volcanic ash. Six elements including Cr, Ni, Cd, Sb, As, and Co constitute heavy metal anomalies. The seismic relics include soft deformation, seismic turbidite, and small graben, etc. The triggering factors for volcanic ash fall and seismic episode were probably the subduction of the Indosian plate beneath the southwestern margin of the South China plate, which formed island arc magmas and volcanic eruptions, meanwhile the back arc basin tension caused earthquakes. Volcanic ash not only provides substances such as aluminium and iron for bauxite-bearing rock series, but also heavy metals which generated to heavy metal anomalies.

**Key words:** catastrophic events; bauxite-bearing rock series; volcanic ash; heavy metals; western Guangxi

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灾变是突变的一种极端形式,多记录于沉积岩中,称为事件地层,与其上下的地层差别显著。灾变事件可分为起因事件、派生事件和终极事件。派生事件在时空上相互叠加或复合,根据其起因类型还可分为地内事件和地外事件,根据时间尺度,又可分为瞬时事件、短时事件和大于10万年的中长期事件。瞬时事件包括地震、海啸、陨石撞击等快速完成的灾变;短时事件,如洪流、火山喷发等几小时到10万年尺度的灾变,反映在毫米-米级厚度的地层中,可再分为沉积事件、化学事件、生物事件和复合事件;大于10万年的中长期事件,如缺氧-富氧、极冷-极热、海进海退、成矿事件、浊流沉积、冰期-间冰期沉积等,呈现出旋回性,是地球层圈耦合过程的响应。内因和果为正反馈关系,如俯冲作用使冷地壳物质沉入地幔,扰动地核外圈,地核外圈迫于热平衡而向地幔补充热物质,进而孕育地幔柱,促发地幔对流,而地幔对流反过来又促进地壳俯冲,两者协同演化,发展到一定程度时,即爆发火成岩省(张克信等,1989; 龚一鸣等, 2007; 殷鸿福等, 2013)。

通过对黑色泥(页)岩、地震岩、海啸岩、微生物岩、重力流沉积岩、风暴岩等进行研究,多数灾变事件得以恢复,构造背景趋于明晰。有些地质灾变遗迹可开辟成旅游景点,如生物群、火山群等;有些衍生为珍贵的物品,如完整的动物化石、陨铁、陨石、文化层等;有些则成为大宗产品,如地震岩和风暴岩被打造成奇石和景观石,黑色生屑灰岩加工成饰面石材,熔岩凝灰岩加工成建材,可见灾变遗迹具相当高的社会效益和经济价值。有些灾变遗迹则可作为

天然实验室,用以观察地质过程,比如重力流、巨厚岩层(泥岩或页岩)软变形、冲击坑、堰塞湖、沥青湖等,科学价值高,所以发现和研究灾变事件具有重要意义。

灾变事件具有等时性、区域性乃至全球性,可作为地层研究的对比界面,如乔秀夫等(1994, 2001)利用地震序列论证了辽东半岛为原地系统,震旦纪之前中朝地台为统一的地体,震旦纪开始到石炭纪早期演化为裂陷槽;杜远生等(2008)通过海相重力流分布推断,在晚泥盆世弗拉期和法门期之交古特提斯海域发生了行星撞击地球事件;余文超等(2013a)研究了铝土矿层的有机物,确定黔北务正道地区铝土矿沉积时期为湿润-干热交替季节;刘平等(2015)从风暴沉积中判别出黔中含铝岩系产出于近海平原中。成岩成矿方面,海啸袭来可瞬间极大推高水柱达30 m,沉积物上覆压力陡增,发生系列短暂的物理、化学反应,如有机质降解、沥青残留、黄铁矿生成、碳酸盐溶解迁移、气体逃逸、矿物快速结晶、沉积物硬化(乔秀夫等, 1989, 1994; 王兆明等, 2006)及缝合线的形成(谭钦银等, 2011);快速沉积可以封存大量的浓缩的海水、膏盐层和有机质,以至于演化为含矿热液(李欣航等, 2023)或盆地卤水,或营造有利的沉淀条件(谢贤洋等, 2016);黄宏伟等(2009)论证了广西丹池地区泥盆系众多事件与成矿的联系,列举了浊积岩、热水沉积岩、滑塌岩、地震岩等4个事件层;杨凯等(2023)认为广西湖润锰矿床中锰质来源受地震事件影响;梁定益(2009)等认为地下热水热气可以软化硬岩石,从而在地震时

发生软变形,成矿流体亦可促成软变形(杜远生等,2017)。黑色页岩更是众所周知的能源矿产,同时也是揭示生物多样性和更替的重要窗口(沈俊等,2010;邱振等,2019)。在国外,有研究者认为地震可引起温度和压力的急剧降低,从而促使Pb-Zn-Cu从热液中沉淀下来(Bozkaya and Banks, 2015),志留纪—石炭纪的喷流型盐水不仅形成欧洲的Pb-Zn-Ba矿床,而且诱发了缺氧和生物集群灭绝(Emsbo, 2017)。火山活动还铸就了珍稀的生物群化石,如燕辽生物群和热河生物群化石(汪筱林等,1999;张向东等,2021)。

在目前全球变暖的背景下,查明灾变事件发生规律,规避灾害事故已成为人类面对的紧迫问题,促使人们更多地研究地史上的灾变事件,因此,灾变事件也是解决问题的抓手之一,是人们提高认识水平,拓展地质学科领域的新途径,长期受到人们的关注。

但是识别灾变事件在地质实践中却显得薄弱,经常被忽视,或引发争议,比如将未变质的沉积岩中的软变形解释为埋深背景下高温高压的变质产物,液化砂脉当作风化的花岗岩,崩落石当作原地沉积成因,等等,事件地层的识别方法仍有待深化。因此,笔者拟展示二叠系含铝岩系中的灾变现象,分析其构造背景,阐释事件成因和意义,以期有助于强化事件领域研究。

## 1 地质概况

桂西地区为右江盆地主体之一,处于右江盆地中心地带(图1)。右江盆地在晚古生代—早三叠世是一个颇为特殊的大陆边缘海(殷鸿福等,1999),加里东运动上升为陆,接受长期剥蚀,于泥盆纪开始海侵,至早泥盆世晚期开始强烈裂陷,形成浅水台地

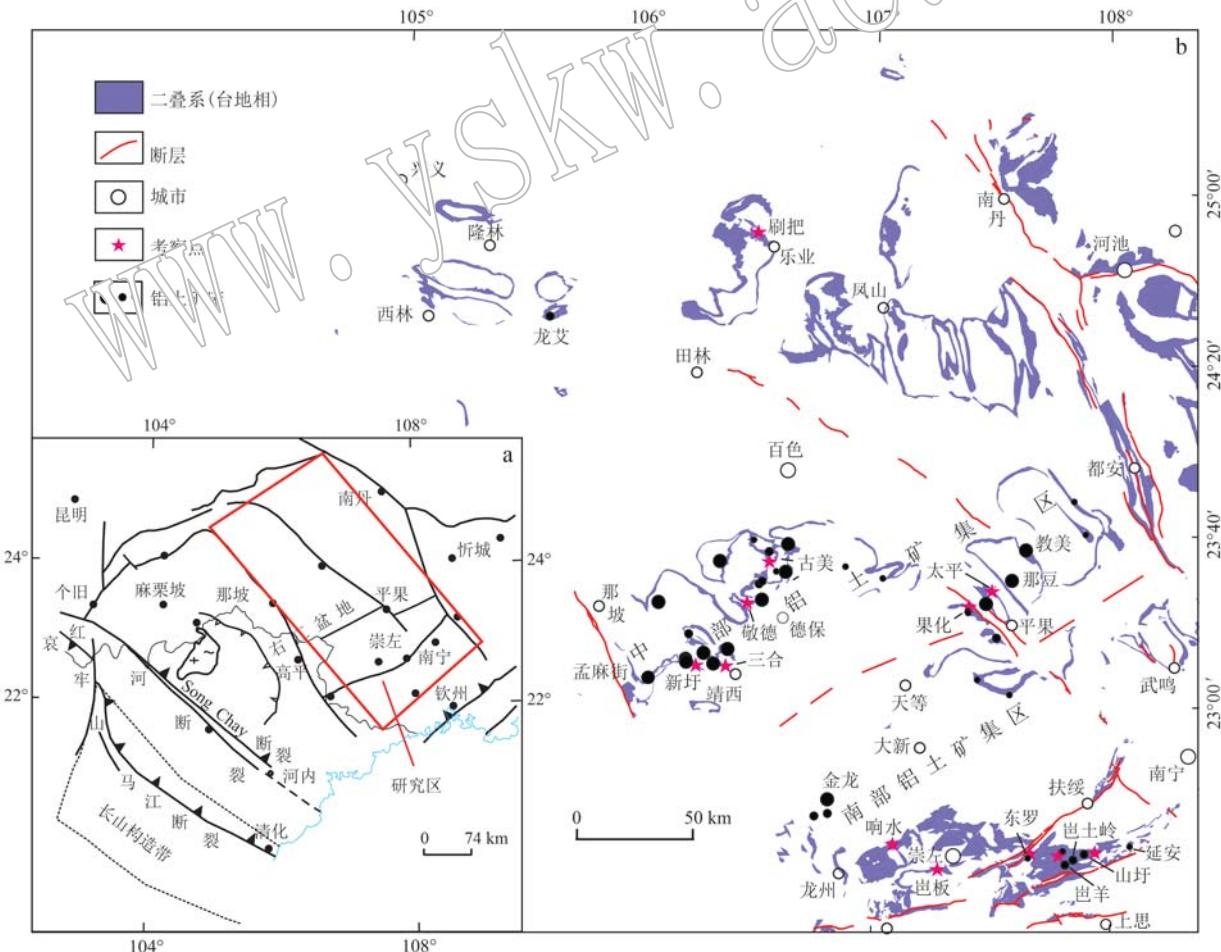


图1 桂西区域构造(a)及二叠系台地与堆积型铝土矿床分布(b)[据钱鑫等(2022)、张启连等(2024a)修改]

Fig. 1 Regional tectonic location (a) and distribution of the Permian platform containing accumulated bauxite deposits (b) of the western Guangxi (modified after Qian Xin et al., 2022; Zhang Qilian et al., 2024)

与台间深水海槽间列的“棋盘式”盆地，此格局具有继承性并保持至早三叠世，发育台地相碳酸盐岩沉积和盆地相碎屑岩沉积（杜远生等，2009），早二叠世在越南境内出现岛弧，右江盆地进入了弧后盆地阶段，直至早三叠世末，岛弧逐渐向东北的中国境内迁移，中三叠世右江盆地成为前陆盆地，沉积了巨厚的浊积沉积建造，中三叠世末，印支地块与华南地块拼合，转入大陆演化至中生代（吴根耀等，2001；杜远生等，2009）。

台地相区（图1中灰色分布区）以浅水沉积为特征，碳酸盐岩主要为生物礁灰岩、泥晶灰岩、泥粒灰岩、鲕粒灰岩，局部白云岩，斜坡则以砾屑灰岩、钙质泥质浊积岩为主，夹火山碎屑岩、硅质岩、泥灰岩，沟槽相或盆地相区以泥岩、硅质岩为主，斜坡相及盆地相均夹有基性-中性火山岩建造（杜远生等，2009）。

泥盆纪末和中二叠世末，右江盆地内碳酸盐台地

一度上升为陆，接受风化剥蚀，沉积含铝岩系，其中上二叠统合山组的含铝岩系分布最广，至第四纪再次风化为堆积型铝土矿（图1）。由于晚二叠世古台地区地形变化较大，以致含铝岩系各微相变化迅速。在出露较为齐全的典型剖面（图2）中可以看到，下部含铝岩系很薄，很少完整显示，以古美ZK1271钻孔揭露较全，为便于叙述，本文将矿层及顶底板称为上含铝岩系，相应地将下部含铝岩系称为下含铝岩系，上、下含铝岩系垂直相距10~15 m，各矿区差异不大，主要矿体赋存于上含铝岩系中，下含铝岩系中矿体规模小。下述的含铝岩系均指上含铝岩系。

上含铝岩系自上而下的完整层序（图3）简述如下。

- ⑥ 炭泥或煤层（线）；
- ⑤ 泥岩；
- ④ 铝土岩夹薄层泥岩；

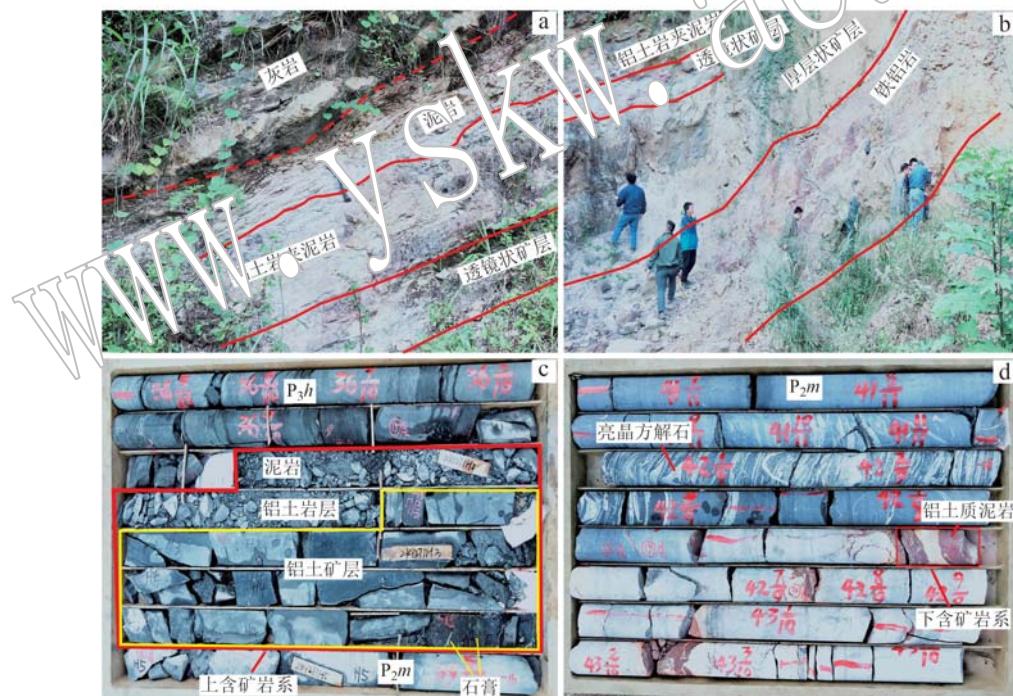


图2 桂西二叠系含铝岩系典型剖面(a、b引自张启连等, 2024b)

Fig. 2 Representative sections of bauxite-bearing rock series of Permian in western Guangxi (a and b from Zhang Qilian et al., 2024b)

a—含铝岩系上部泥岩段，发育波状层理，缺失顶部炭泥和煤层线，平果太平；b—含铝岩系中下部，底部缺失部分铁铝质岩及古土壤，平果太平；c、d—合山组底部的上含铝岩系与茅口组上部的下含铝岩系，前者产出厚层状矿层及铝土岩层，后者仅见铝土岩层，至少两期暴露，田阳古美ZK1271，岩心箱长度1 m

a—mudstone with wavy bedding in the upper portion of bauxite-bearing rock series that lack the carbonaceous mud or coal seam in Taiping of Pinguo；b—middle to lower portion bauxite-bearing rock series in the Taiping of Pinguo, without ferro-aluminous and paleosol at the bottom；c and d—drill core of ZK1271 of Gumei deposit in Tianyang County showing upper bauxite-bearing rock series with chick beds of bauxite and bauxitic rocks at the bottom of the Heshan Formation and lower bauxite-bearing rock series with bauxitic rocks in the upper portion of the Maokou Formation, indicating at least two stages of exposure, length of the drill-core box is 1 m

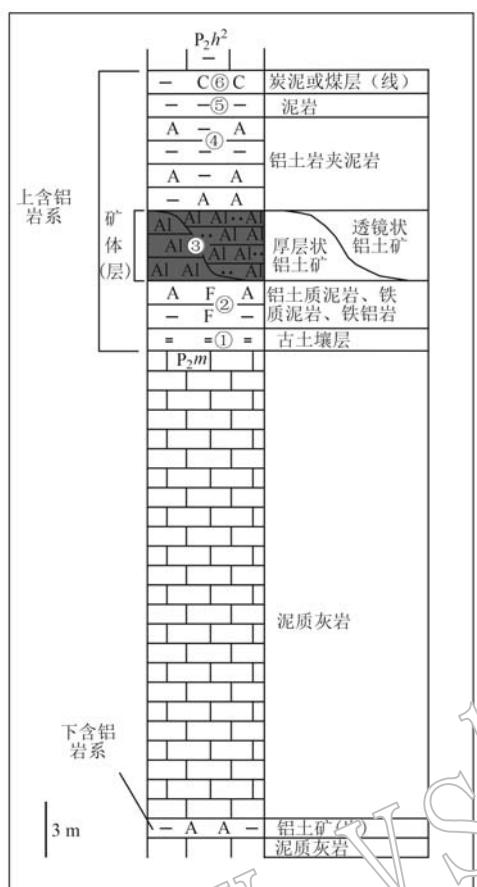


图3 桂西二叠系含铝岩系分布特征

Fig. 3 Characteristics of aluminium-bearing rock series  
of Permian in western Guangxi

- ③ 铝土矿层;
- ② 铁铝质岩;
- ① 古土壤。

铝土矿层可分为厚层状和透镜状(图2a、2b),前者为风化壳的残留,呈板状,块状层理,矿石类型以块状铝土矿石为主,有时见弱定向构造,基质为隐晶-半隐晶,粒屑以鲕(豆)和火山灰为主;透镜状矿层则为凹地沉积,发育小型交错层理、平行层理、韵律层理、粒序层理等,时见块状层理,碎屑状矿石为主,定向构造发育,粒屑多有磨损。

铁铝质岩和黏土质岩沉积并向下过渡到基岩的组合属于典型的残积类型(杨伟东等,1991)。铝土矿中B含量 $10\times10^{-6}\sim20\times10^{-6}$ ,属于淡水环境(张启连等,2016)。

下含铝岩系规模相对较小,地表极少见及。露头尺度上可见岩性从凹地中心的碎屑沉积向边部泥质沉积过渡,过去勘查钻孔略为穿越上含铝岩系即

终孔,故而研究程度较低,据目前有限资料判断,其含矿性差。

## 2 样品采集与测试

样品采集覆盖桂西地区,尽量采全岩(矿)石类型,具体见图4。样品送至澳实分析检测(广州)有限公司检测主、微量元素,破碎后缩分出300 g,研磨至75  $\mu\text{m}$ (200目),采用X射线荧光光谱仪熔融法分析常量元素,检出限为0.01%,电感耦合等离子体质谱法测试稀土和微量元素,分析精度为5%~10%。本文选择讨论的数据为部分重金属元素和稀土元素。

SEM 观察在中南大学冶金学院扫描电镜实验室进行,使用仪器为岛津 E'PMA1720H 束流强度10 nA,束斑直径2  $\mu\text{m}$ 。

岩矿标本送至河北省测绘院实验室磨制成无玻璃光薄片,抛光处理,由笔者完成岩矿鉴定。

## 3 灾变事件

### 3.1 火山灰季

火山灰不仅大量存在于含铝岩系各个分层中(侯莹玲,2017;程顺波等,2021;张启连等,2022),而且顶板海侵的砂泥岩和泥灰岩中亦发现火山灰纹层或薄层凝灰岩。前人铝土矿锆石测年结果表明,数据变化较大,如Deng等(2010)获得2个铝土矿加权平均年龄,分别为 $256\pm2$  Ma(德保)和 $261\pm2$  Ma(平果);Yu等(2016)获得4件铝土矿加权平均年龄,即 $263.0\pm1.2$  Ma(靖西)、 $263.4\pm1.0$  Ma(德保县)、 $262.6\pm1.2$  Ma(扶绥)和 $262.5\pm1.1$  Ma(乐业县);侯莹玲等(2014)在平果获得铝土矿峰值年龄为262 Ma,而铝土矿顶板碎屑岩峰值年龄253 Ma,铝土矿与顶板碎屑岩锆石年龄差值最大达10 Ma,考虑到铝土矿层底部的铁铝质岩和古土壤尚无锆石年龄数据,本文认为火山灰沉降时长可以达到10 Ma,是一个较长的时间段。

含铝岩系上部铝土岩-泥岩、中部铝土矿层及下部铁铝质岩古土壤中的火山灰展示于图5。上部铝土岩-泥岩岩层中的火山灰中常见石英、碱性长石、斜长石等代表酸性岩浆的斑晶(常丽华等,2009)(图5a~5c)。桂西北部乐业县刷把村附近泥质铝土岩中的球泡,外壳厚度近乎一致,具梳状结构,壳与内部成分相似,个别球泡外壳呈港湾状,呈现塑性

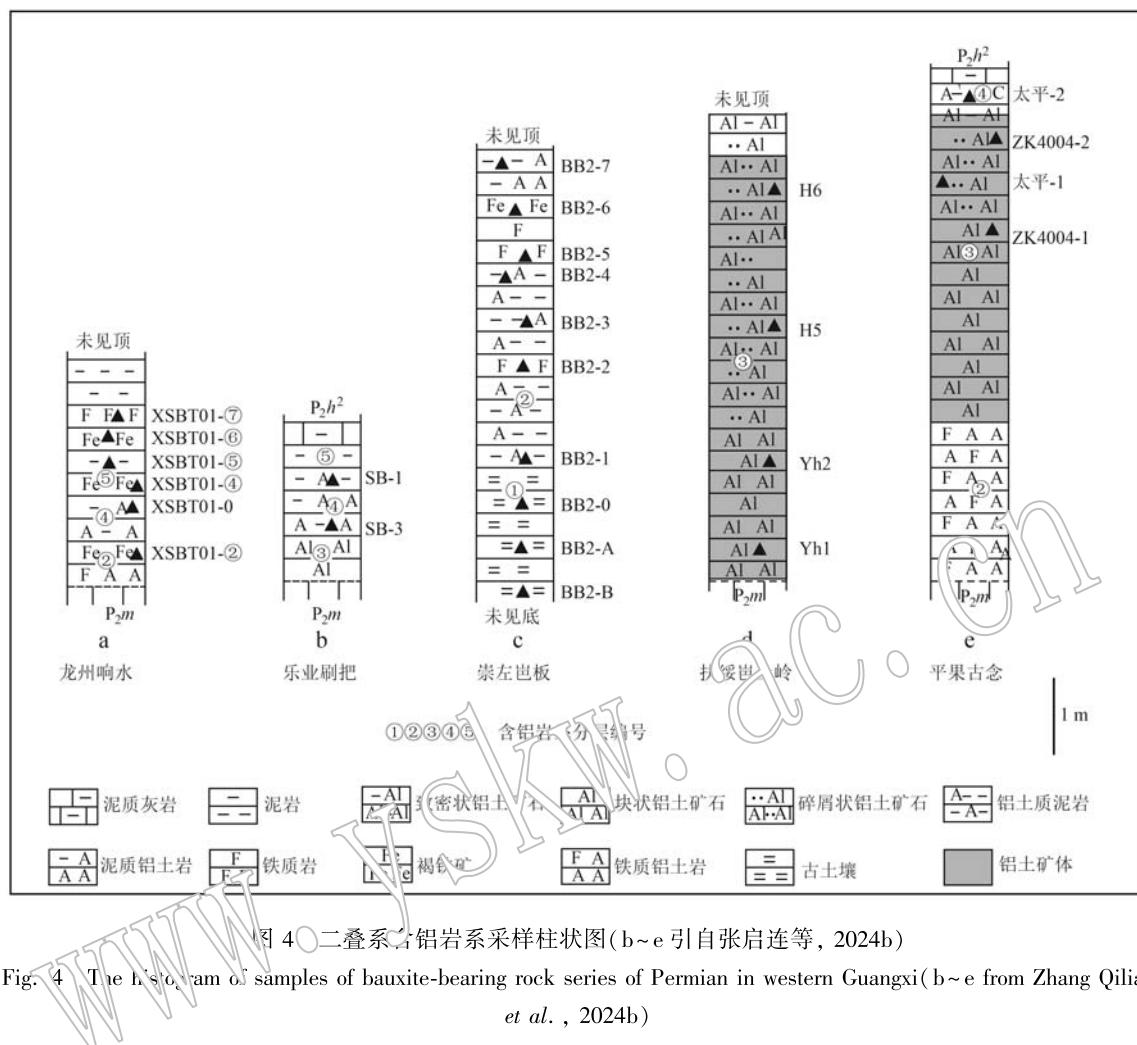


图 4 二叠系含铝岩系采样柱状图(b~e 引自张启连等, 2024b)

Fig. 4 The histogram of samples of bauxite-bearing rock series of Permian in western Guangxi (b~e from Zhang Qilian et al., 2024b)

变形(图 5a),球泡被认为是黏度很大的酸性熔岩的产物。桂西中部平果太平铝土矿段泥质铝土矿石(致密状矿石)中的火山灰中,石英斑晶被熔蚀,浅灰色塑性岩屑已轻微脱玻化,极细粒晶核尚可分辨(图 5b),浅色玻璃暗示其为流纹质岩浆(常丽华等, 2009)。南部扶绥喀兰铝土质泥岩中的石英、碱性长石、斜长石晶屑具有熔蚀结构和碎斑结构(图 5c)。中部厚层状铝土矿层中的火山灰,尽管经历了强烈淋滤,但仍可分辨其中具有壳的浆屑或塑性玻屑、结构均一的玻屑以及边角熔蚀的辉石晶屑(5d~5f)。下部铁铝质岩或古土壤中火山灰保存最好,镜下几乎全为火山灰,安山岩、玄武岩屑常见(图 5g~5j),安山岩中长石多呈微晶状,略显定向,另有少量玻璃,局部见球粒结构,其中斜长石的中空结构清晰,玄武岩屑中斜长石晶体稍粗大,杂乱不显定向,基质中铁质(黑色)居多,未见石英;上部铁质铝土岩中还见到橄榄石假晶,文象构造尚存,个别晶屑具有反

应边,基质以玻璃为主,浅灰色者应为磁铁矿。

自下而上,斑晶从辉石到石英斑晶,火山灰从玄武质、安山质变化为流纹质,具有钙碱性岩浆特点。铝土矿中尚有一些特殊的矿物,如铬铁矿、自形锆石(图 6),表明有基性-超基性岩浆参与了火山喷发。

### 3.2 重金属元素异常

重金属元素包括 Hg、Cd、Pb、As、Cr、Co、Ni、Cu、Zn 等,在土壤形成的重金属污染通常被认为受成土母岩主导(邢润华等, 2022),现代土壤中这些元素富集与铁锰结核数量呈正相关关系(杨琼等, 2021)。表 1 为桂西上二叠统合山组含铝岩系中除 Hg 之外的重金属含量,除 Pb、Zn 富集系数接近 1.5 外,其他 6 个元素均有一定的富集。同为铝土矿富集区的黔北务正道地区,铝土矿及黏土岩相关元素含量并不高,如 Cu 最高值为  $20.8 \times 10^{-6}$ , Cr 最高值为  $456 \times 10^{-6}$ , Co 最高值为  $54 \times 10^{-6}$ , Ni 最高值为  $101 \times 10^{-6}$ , Cd、Sb、As 未见数据(张莹华等, 2013),华

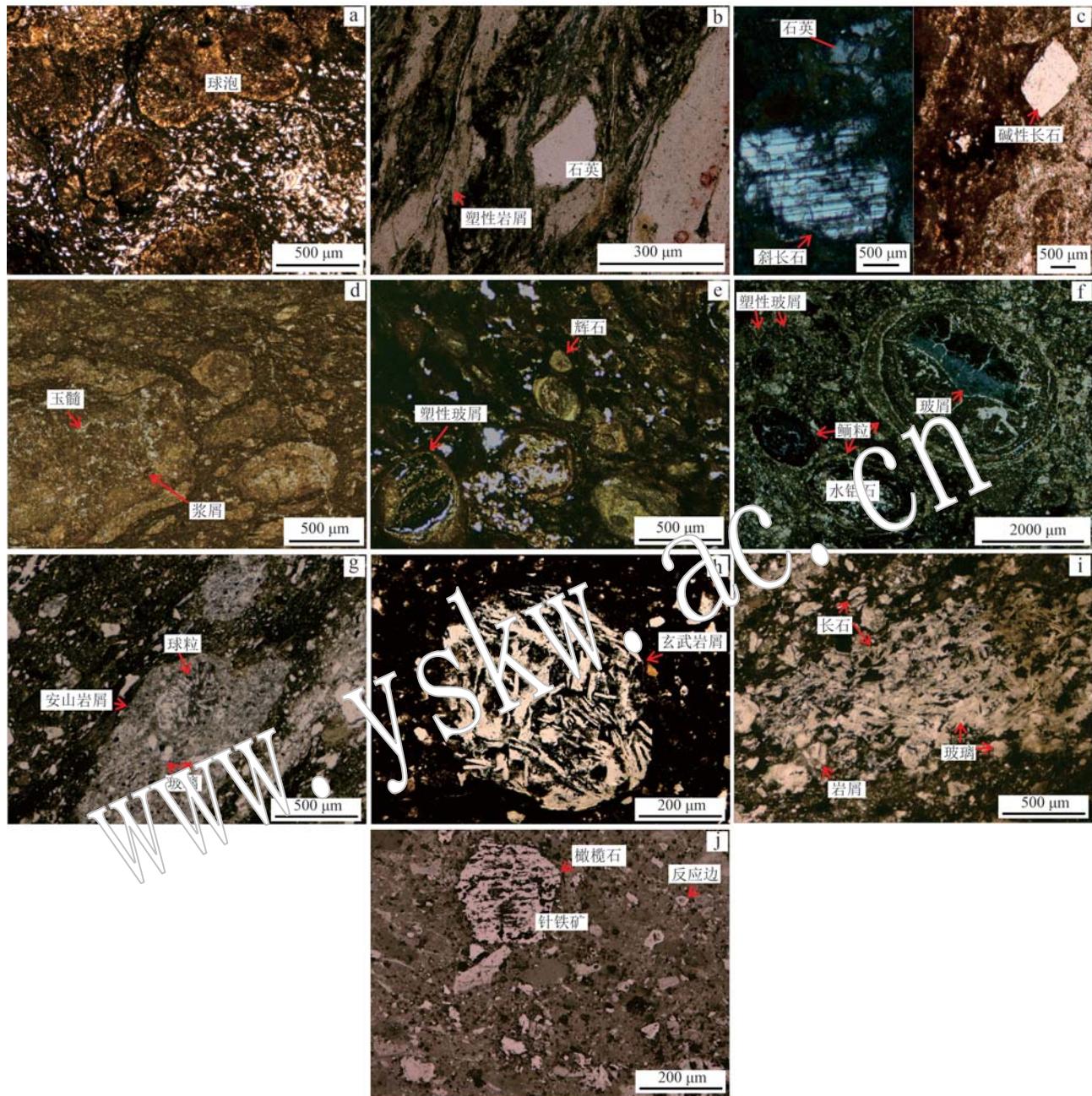


图 5 桂西二叠系含铝岩系典型火山灰特征

Fig. 5 Typical characteristics of volcanic ash in the Permian bauxite-bearing rock series in western Guangxi

a—致密状铝土矿中的球泡,具壳,呈梳状构造,酸性火山灰,乐业,SB-2; b—铝土质泥岩中的塑性岩屑和石英斑晶,酸性火山灰,平果,太平-2; c—铝土质泥岩中的石英、碱性长石、斜长石斑晶,具熔蚀结构,扶绥,KLSK-1; d—块状铝土矿石中的浆屑,内部局部脱玻,具壳,平果,福布1; e—块状铝土矿石中的塑性玻屑和晶屑,扶绥,渠坎1; f—块状铝土矿石中的塑性玻屑,靖西,新圩X; g—铝土质泥岩中的安山岩屑,基质中长石微晶略呈定向,可见球粒,崇左,BB2-1; h—铁铝岩中的玄武岩屑,浑圆状;崇左,BB2-7; i—铁质铝土岩中的安山岩屑,斜长石呈中空状,玻基交织结构,崇左,BB2-2; j—铁质铝土岩中的橄榄石假晶,针铁矿在晶屑中呈星象结构,部分晶屑具反应边,基质中针铁矿大小悬殊,呈浸染状,针铁矿原矿物为磁铁矿,崇左,BB2-2,反射光

a—the lithophysae in pelitic bauxite, with comb-like crust, presenting acidic magma, Leye County, SB-2; b—the plastic detritus and quartz phenocryst presenting acidic magma in aluminaceous mudstone, Pingguo City, Taiping-2; c—the porphyry of quartz, alkaline feldspar, plagioclase in aluminaceous mudstone, showing resorption texture, Fuishui County, KLSK-1; d—the magma fragment in massive ore, showing inner devitrification and a shell, Pingguo city, Fubu 1; e—the plastic detritus and crystal fragment in massive ore, Fushui County, Qukan1; f—the plastic hyaloclastite fragment in massive ore, Jingxi City, Xinxu X; g—andesite debris in aluminaceous mudstone, with feldspar microcrystals are slightly oriented in the matrix, and a spherule displaying, Chongzuo, BB2-1; h—round basaltic debris in ferroaluminite rock, Chongzuo, BB2-7; i—andesite debris in ferruginous bauxite, with hyalopilitic texture in plagioclase, Chongzuo city, BB2-2; j—the olivine pseudocrystal in ferruginous bauxite, showing goethite (light gray) constituting the graphic texture, some of the crystal fragments circled by reacting edge, the size of goethite in the matrix vary wildly and disseminated, the primary mineral of goethite was magnetite, Chongzuo City, BB2-2, reflecting light

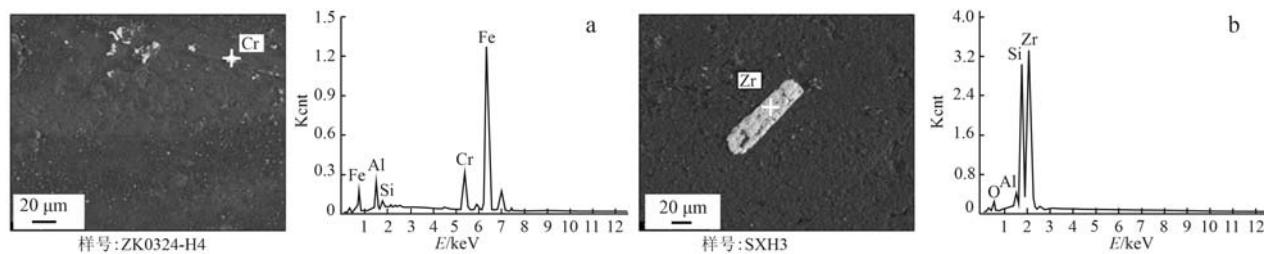


图 6 桂西上二叠统含铝岩系中铬铁矿和锆石的矿物化学成分

Fig. 6 Composition of chromite and zircon of in bauxite-bearing rock series of the upper Permian in western Guangxi

a—铝土矿中的铬铁矿, 扶绥山坪; b—铝土矿中的自形锆石, 扶绥山坪

a—chromite in bauxite of the Shanxu of Fushui County; b—euhedral zircon in bauxite of the Shanxu of Fushui County

表 1 桂西上二叠统含铝岩系部分重金属元素含量

 $w_B/10^{-6}$ 

Table 1 The contents of heavy metal in bauxite-bearing rock series of Upper Permian in western Guangxi

剖面	样号	岩石类型	Cd	Pb	Cu	Co	Ni	Cr	S	In	Sb
龙州响水	XSBT01-⑦	褐铁矿	2.16	23.7	35.7	68.4	302	1 150	9.3	10	0.89
	XSBT01-⑥	铁铝岩	0.79	29.3	54.8	210	35	780	16.2	152	1.98
	XSBT01-⑤	铁铝岩	2.05	27.2	44.7	71.7	31	1 130	9.9	111	1.06
	XSBT01-④	褐铁矿	1.31	27.7	52.8	183	338	530	16.8	126	1.96
	XSBT01-0	铝土岩	4.28	9.1	57.1	170.5	600	780	2.5	120	0.41
	XSBT01-②	褐铁矿	0.84	9.6	36.4	224.0	361	740	10.3	149	1.39
平均			1.91	22.7	40.9	14.68	379.67	851.67	10.83	127.33	1.28
富集系数			24.6	1.3	1.24	4.83	6.66	13.52	5.70	1.48	8.54
乐业刷把	SB-3	泥质铝土岩	0.22	5.3*	454*	45.0*	68.7*	600*	9.5	171*	2.82
	SB-1	铝土质泥岩	0.06	1.8*	369*	17.2*	92.1*	770*	5.4	161*	1.90
	平均		0.14	5.55	411.50	31.10	80.40	685.00	7.45	166.00	2.36
	富集系数		0.55	0.37	10.83	0.97	1.41	10.87	3.92	1.93	15.73
崇左岜板	BB2-7	铝土质泥岩	0.28	23.4*	56.2*	26.5*	333*	650*	7.9	54*	1.15
	BB2-6	泥质铝土岩	0.11	26.7*	76.7*	49.9*	352*	1 000*	9.0	113*	1.51
	BB2-5	铝土质泥岩	0.33	27.8*	22.0*	17.9*	322*	600*	6.1	36*	1.25
	BB2-4	铝土质泥岩	0.11	6.2*	18.5*	18.9*	452*	740*	2.9	30*	0.26
	BB2-3	铝土质泥岩	0.18	21.5*	32.9*	20.9*	394*	1 020*	7.6	35*	0.92
	BB2-2	铝土质泥岩	0.06	19.7*	100.0*	18.8*	460*	830*	5.2	72*	0.70
	BB2-1	铝土质泥岩	0.12	10.2*	228*	28.2*	383*	1 430*	3.9	78*	0.23
	BB2-0	古土壤	0.12	4.8*	886*	166.5*	560*	780*	2.7	268*	0.09
	BB2-A	古土壤	0.33	13.4*	184.5*	20.3*	380*	1 440*	4.1	97*	0.36
	BB2-B	古土壤	0.11	7.8*	264*	73.3*	512*	1 250*	12.5	90*	0.27
	平均		0.18	16.15	186.88	44.12	414.80	974.00	6.19	87.30	0.67
	富集系数		3.18	1.08	4.92	1.38	7.28	15.46	3.26	1.02	4.49
扶绥岜土岭	H6	碎屑状铝土矿	0.35	45.5*	11.9*	4.7*	36.9*	660*	13.4	19*	1.43
	H5	碎屑状铝土矿	0.23	48.9*	10.1*	3.6*	34.5*	470*	4.6	16*	1.56
	YH2	块状铝土矿	0.31	75.4*	3.5*	22.7*	111.0*	630*	12.8	18*	1.94
	YH1	块状铝土矿	1.34	64.1*	3.2*	19.5*	113.0*	820*	30.4	20*	2.71
	平均		0.56	58.48	7.18	12.63	73.85	645.00	15.30	18.25	1.91
	富集系数		10.14	3.90	0.19	0.39	1.30	10.24	8.05	0.21	12.73
平果太平	ZK4004-2	致密状铝土矿	<0.02	21.5*	10.5*	2.2*	22.3*	100*	11.0	4*	3.00
	ZK4004-1	块状铝土矿	0.09	58.8*	15.6*	3.6*	28.0*	950*	38.2	4*	7.43
	太平-2	致密状铝土矿	0.06	8.4*	5.9*	5.2*	46.7*	60*	15.9	5*	2.11
	太平-1	碎屑状铝土矿	0.16	47.7*	6.3*	9.1*	20.8*	520*	16.9	12*	4.66
	平均		0.10	34.10	9.58	5.03	29.45	407.50	20.50	6.25	4.30
	富集系数		1.88	2.27	0.25	0.16	0.52	6.47	10.79	0.07	28.67
中国陆壳			0.06	15.0	38.0	32.0	57.0	63	1.9	86.0	0.15

\*—引自张启连等(2024b); 中国陆壳数据来自黎彤(1994)。

北“G”层铝土矿及黏土岩近年研究认为火山灰参与了成矿(Liu et al., 2014; 刘学飞等, 2020; 张保涛等, 2023),山西石墙铝土矿区Cu含量一般小于 $100\times10^{-6}$ ,个别特高值为 $1228.8\times10^{-6}$ ,Ni含量一般为 $100\sim300\times10^{-6}$ ,个别特高值为 $509\times10^{-6}$ ,Co含量一般 $10\sim100\times10^{-6}$ ,最高值为 $173.1\times10^{-6}$ ,Cd、Cr、Sb、As未见数据(孙思磊等, 2012)。在国外,除希腊的Parnassos-Ghiona略高于本区外,已报道的喀斯特型铝土矿区其Cr、Ni含量均低于桂西(Laskou and Economou-Eliopoulos, 2007)。桂西地区含铝岩系的Cr、Ni为强烈富集,Cr在所有的剖面中均具有高的

含量,平均值 $400\times10^{-6}\sim900\times10^{-6}$ ,Ni含量在部分剖面如响水和岜板剖面亦较高,一般为 $300\times10^{-6}\sim600\times10^{-6}$ ,Cd为富集(富集系数 $1.5\sim5$ )至强烈富集(富集系数 $\geq 5$ ),Sb、As则在酸性火山灰为主的样品中强烈富集,如平果太平剖面,Sb一般含量为 $1\sim10\times10^{-6}$ ,As一般含量为 $10\sim40\times10^{-6}$ ,Co局部富集,Cd、Cr、Ni、Co、Sb、As这6个元素构成了重金属异常。

### 3.3 地震幕

含铝岩系中的古地震遗迹时有见及,其中软变形和震浊积岩保存较好(图7)。

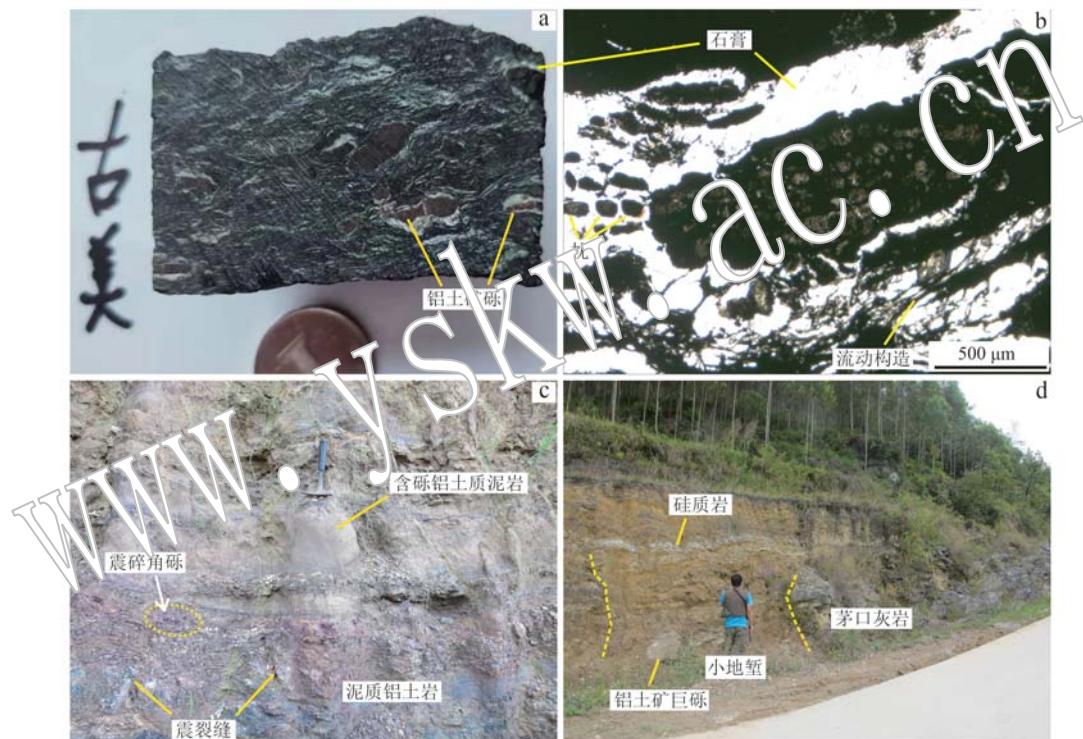


图7 桂西二叠系含铝岩系中古地震遗迹

Fig. 7 Relics of Permian earthquakes in bauxite-bearing rock series in western Guangxi

a—铝土矿石中的石膏(白)构成流动构造,裹挟铝土矿或铁铝岩砾,田阳古美; b—石膏(白色)形成微观流动构造,将铝土矿分割成枕,田阳古美; c—含铝岩系顶板的震浊积岩,A段为震裂角砾岩,地裂缝未穿越上覆岩层,铝土质泥岩中有“漂砾”,德保马隘; d—小地堑,内有铝土矿巨砾,沉积铝土质泥岩,上覆沉积为泥岩夹硅质岩,层理完整,德保马隘

a—gypsum (white) in the bauxite ore formed a flowing structure, containing bauxitic or ferroaluminite boulders, Gumei of Tianyang; b—gypsum (white) formed microfluidic structure that splits the bauxite into pillows, Gumei of Tianyang; c—seismic turbidite lying over bauxite-bearing rock series, showing the A park like shattered rock, and the ground fissures not penetrating the overlying layer, while the boulder scattering within bauxite mudstone, Ma'ai of Debao County; d—a small graben filled giant boulders of bauxite, covered by mudstone which wrap successive siliceous rock, Ma'ai of Debao County

软变形(图7a、7b),见于田阳古美矿区ZK1111,发育于厚层状铝土矿中,共有3段,厚20~50 m,每段的顶、底板为深灰色铝土矿,基本不变形,而含石膏的矿层发生了流动,可排除滑塌变形成因,铁铝岩

砾呈叠瓦状或旋转,或被石膏穿刺,镜下石膏呈白色,局部形成流动构造,有时将铝土矿砾分割呈枕状,推测是地震时充水,石膏吸水后在主震期发生液化,顺层流动过程中对铁铝岩砾和深灰色铝土矿进

行穿刺破裂甚至挟裹其移动所致。

震浊积岩(图7c),由含铝岩系上部铝土质泥岩构成,与经典的浊积岩一样具有自下向上变细的递变层理,即鲍马层序,但震浊积岩的鲍马层序AB段以含震裂砾石为特点,发育不穿透上覆岩层的震裂缝或塑性变形构造,岩层不对应,不协调(赵卫卫等,2006)。较厚的泥岩段可见悬浮状砾石弥漫状分布,说明紊流较强,属于浊流沉积(李林等,2011)或阵发式浊流(曾允孚等,1986)。

小地堑(图7d),见于德保马隘矿区,地堑一壁为灰岩,另一壁为铝土质泥岩夹铝土岩,内部沉积铝土质泥岩,层理与两侧岩层不相连,接触处较乱不显层理,可见铝土矿巨砾,与周围具层理的铝土质泥岩不协调,上覆铝土质泥岩夹硅质岩,层理连续性好,属海相沉积,灰岩一肩位的含铝岩系变薄。和上述震浊积岩一样,小地堑形成于含铝岩系上部铝土质泥岩沉积阶段,是古地震的特征之一(梁定益等,1991)。

除了上述4种特殊现象外,在三合矿区含铝岩系顶板泥灰岩中见到球-枕构造和泄水孔。在下含铝岩系中见到铝土质泥岩呈液化砂脉形式侵入到上覆灰岩中,脉体呈现下宽上窄特征。在含铝岩系及其顶底板均有地震带,表明有较密集的多期次地震发生并成一个地震带(乔秀夫等,2009)。

## 4 讨论

### 4.1 弧火山喷发和弧后盆地扩张是灾变事件的诱因

多数研究者通过海相火山岩研究认为晚二叠世右江盆地为弧后盆地(陈洪德等,1990;张锦泉等,1994;曾允孚等,1995;杜远生等,2009),董云鹏等(1999)在滇东南亦发现了石炭纪一二叠纪岛弧火山岩,胡丽沙等(2012)在那坡发现了早中三叠世岛弧火山岩,证明了弧后盆地在早二叠世就已出现。笔者在桂西二叠系含铝岩系中发现安山岩岩屑与玄武岩岩屑除大量分布于岩系底部的铁铝质岩和古土壤中外,至岩系上部铝土质泥岩中仍有出现;安山岩岩屑中长石呈微晶状,具中空结构,经多个薄片考察,尚未见其作为斑晶产出,表明喷发岩浆房富水,抑制了长石结晶而富集在喷发熔浆中(Wang et al., 2022),导致喷发时快速结晶(图5g),最有可能是俯冲板片析出了流体使源岩区富水(高俊等,2024);大量磁铁矿(已转化为针铁矿)出现表明铁趋向于后

期富集,并且后期同时出现碱性长石、斜长石、石英斑晶共生(图5e)显示流纹质岩属性,玄武岩-安山岩-流纹岩组合指示岩浆源为钙碱性系列(常丽华等,2009;刘阁等,2012)。本次工作发现含铝岩系自上至下Cr、Ni含量均较高,表明源岩浆房物质主要来自地幔楔,被俯冲板片析出流体交代熔融,体现出岛弧岩浆中微量元素典型的地球化学特征,如富集轻稀土元素(L/H平均值3.4)和某些大离子亲石元素(Th、U、Pb),亏损高强场元素(Nb、Ta)(张启连等,2022),证明本区铁铝质岩沉积时期桂西邻区已有成熟的俯冲带存在,而整个含铝岩系均有高含量的Cr、Ni出现表明有过多次俯冲,火山的喷发是俯冲作用最强烈的表现,是高角度俯冲作用的产物(李政林等,2019;肖文交等,2022)。火山灰中岩屑的存在表明火山灰源区不远,如现代日本列岛火山灰在100~600 km的范围内(陈宣渝等,2014, 2022),所以从距离上考虑,印支造山带岛弧火山喷发与本区距离适中且晚二叠世印支板块和华南板块之间的岛弧正处在鼎盛时期(Jian et al., 2009; Shi et al., 2015; 林伟等, 2024),所以它应是桂西火山灰的来源地。现代全球最主要的火山喷发和地震均沿着环太平洋和新特提斯造山带分布,在岩浆弧区,俯冲板片释放流体交代地幔楔、板片断离、后撤、弧后盆地扩张等各种构造活动,均可引发火山和地震的发生(肖文交等,2022),桂西当时构造位置为弧后盆地,推断弧火山和弧后盆地扩张是本区火山灰沉积、重金属污染和地震的诱因。

### 4.2 与铝土矿成矿作用的关联

现代近赤道红土型铝土矿均与岩浆岩风化壳有关,岩浆岩中较高的Al含量可使原岩经历一次风化淋滤即可形成工业矿石;而以碳酸盐岩为基岩的风化壳需要再次淋滤才能形成铝矿石,如黔北和山西地区的二叠、石炭系铝土矿(甄秉钱等,1986;王银川等,2011;余文超等,2013b),桂中现代红土型铝土矿直接与古生代碳酸盐岩接触,虽然已形成瘤状、豆瓣状构造,但仍未达到工业利用价值(赵辛金等,2021),表明源岩高Al含量是形成铝土矿的重要因素。桂西地区二叠系含铝岩系中多次淋滤迹象不明显,大多矿区中仅少数出现,厚度仅数厘米,沉积后再遭受淋滤的机制并未如黔北和山西的铝土矿那样普遍。显然,桂西中晚二叠世的铝土矿中铝的富集机制,不在于多次淋滤,而是依赖高铝的物质即火山灰的参与。

与火山灰缺失或较少的黔北铝土岩系和华北铝土岩系相比,桂西铝土岩系中 Cd、Cr、Co、Ni、Sb、As 含量均较高,且具有富集至强富集特点,显然是火山灰带来的结果。本区发现针铁矿组成的文象结构表明,火山灰降落事件对 Fe 富集同样有利。上述 6 个元素富集与含铝岩系中铬铁矿及金、铂族矿物共存(余文超, 2017),表征超基性-基性岩浆喷发。最近研究表明,半金属元素 As、Bi、Te、Se、Sb 易与铂族元素结合为纳米级合金,它们之间的亲和力强于铂族元素与硫化物的亲和力,且能在热液中迁移(王焰等, 2023)。既然火山灰带来的重金属和稀贵金属可在台地中心含铝岩系中富集,亦可在相邻海域中富集,除了富集重金属和稀贵金属,海水亦可富集于重力流快速沉积中,因为快速沉积既能封存巨量的海水、有机质,亦可封存深部来源的热水;封存水大有可能在后期的构造活动中演化为热卤水,萃取矿质并向低压区运移,从而成矿。滇黔桂地区微细粒型金矿的古流体年龄为 260 Ma 左右,与含铝岩系沉积年代相当,其中硫主要具海水硫酸盐同位素特征,水则为地层水、大气降水和岩浆水的混合成因(胡清宽等, 1995),有可能与该时期火山和地震事件有关。

## 5 结论

(1) 桂西二叠系含铝岩系主要灾变事件为火山灰季、重金属异常和地震幕。火山灰降落时长可达 10 Ma, 火山灰岩性为基性到酸性均有出现, 属岛弧钙碱性岩浆系列; 具有 Cd、Cr、Co、Ni、Sb、As 等重金属异常; 古地震遗迹见有软变形、震浊积岩、小地堑等。

(2) 火山灰降落和地震幕的诱因为印支板块和华南板块西南缘之间的俯冲, 汇聚发展到高角度俯冲时激发岛弧岩浆喷发, 同时弧后盆地的拉张活动引发了地震。

(3) 火山灰不仅有利于铝土矿、铁矿的成矿, 还加强了重金属的富集。

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